

Article

Radical-Driven Methane Formation in Humans Evidenced by Exogenous Isotope-Labeled DMSO and Methionine

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Abstract: Methane (CH₄), produced endogenously in animals and plants, has recently been suggested to play a role in cellular physiology, potentially influencing the signaling pathways and regulatory mechanisms involved in nitrosative and oxidative stress responses. In addition, it has been proposed that supplementation of CH₄ to organisms may be beneficial for the treatment of several diseases, including ischemia, reperfusion injury, and inflammation. However, it is still unclear if and how CH₄ is produced in mammalian cells without the help of microorganisms, and how CH₄ might be involved in physiological processes in humans. In this study, we provide first proof of the principle that CH₄ is formed non-microbially in the human body by applying isotopically labeled methylated sulfur compounds, such as dimethyl sulfoxide (DMSO) and methionine, as carbon precursors to confirm cellular CH₄ formation. A volunteer applied isotopically labeled (²H and ¹³C) DMSO on the skin, orally, and to blood samples. Monitoring of stable isotope values of CH₄ convincingly showed the conversion of the methyl groups, as isotopically labeled CH₄ was formed during all experiments. Based on these results, we consider several hypotheses about endogenously formed CH₄ in humans, including physiological aspects and stress responses involving reactive oxygen species (ROS). While further and broader validation studies are needed, the results may unambiguously serve as a proof of concept for the endogenous formation of CH₄ in humans by a radical-driven process. Furthermore, these results might encourage follow up studies to decipher the potential physiological role of CH₄ and its bioactivity in humans in more detail. Of particular importance is the potential to monitor CH₄ as an oxidative stress biomarker if the observed large variability of CH₄ in breath air is an indicator for physiological stress responses and immune reactions. Finally, the potential role of DMSO as a radical scavenger to counteract oxidative stress caused by ROS might be considered in the health sciences. DMSO has already been investigated for many years, but its potential positive role for medical use remained highly uncertain.

Keywords: endogenously produced methane; reactive oxygen species; methyl radicals; dimethyl sulfoxide; methyl donors; stable carbon and hydrogen isotope labeling; oxidative stress; antioxidant defense system

1. Introduction

Methane (CH₄) is an important and highly abundant carbon molecule in the Earth's atmosphere that affects the Earth's radiative balance. Around 600 to 700 million tons of CH₄ are released annually to the atmosphere by natural and anthropogenic sources, mostly of biological origin [1]. For a long time, biological CH₄ formation was considered to only occur from the metabolism of microorganisms—methanogens that belong to the domain Archaea—living under strictly anaerobic conditions in natural wetlands, landfills, rice fields, or in the alimentary tract of vertebrates, including

ruminants or humans. However, in 2006 it was clearly demonstrated that plants—from the Eukarya domain—are able to produce CH₄ per se [2]. This breakthrough finding fueled research to search for alternative biological CH₄ sources other than those derived from archaeal methanogens. Since then, many studies have unambiguously confirmed direct (endogenous) CH₄ formation and release from eukaryotes, including plants [3-8], animals [9, 10], fungi [11, 12], and marine and freshwater algae [13-15]. In addition, cyanobacteria—belonging to the domain Bacteria—thriving in aquatic and terrestrial environments are also able to generate CH₄ at substantial rates depending on species and environmental conditions [16]. Finally, a universal mechanistic formation for CH₄ has recently been proposed that might explain why all living organisms produce CH₄ under aerobic conditions [17]. In conclusion, the recent findings have stimulated great interest in endogenous CH₄ formation processes all over the three life-domains, which are now often referred to as “aerobic”, “non-archaeal”, or “non-microbial” CH₄ formation processes [18]. Therefore, in the following paragraphs we first discuss the traditional view of CH₄ formation before we deal with the recently identified alternative pathways.

Traditional view of microbial methane formation in humans

Traditionally, CH₄ formation in humans was considered to exclusively arise from anaerobic microbial activity in the gastrointestinal tract [19-21]. The microbial species identified so far in the distal part of the colon (methylotrophic *Methanospaera stadtmaniae* and hydrogenotrophic *Methanobrevibacter smithii*) were considered to contribute to the observed 25% to 70% of humans defined as CH₄ producers. The terms “CH₄ producer” (with breath CH₄ emission > 1 part per million (ppmv) above background values) and “CH₄ non-producer” (breath CH₄ emission < 1 ppmv above background values), used in many previous publications, have become somewhat misleading after Keppler et al. [22] demonstrated through high precision measurements that every human produces breath CH₄ at least 26 parts per billion by volume (ppbv) above background CH₄ values. This value is around three orders of magnitude lower if compared with the average values from typical breath CH₄ “producers”. Based on these results, it was proposed to preferentially use the terminology of high (> 4 ppmv), medium (1–4 ppmv), and low (< 1 ppmv) breath CH₄ producers instead of CH₄ producers and non-producers.

Regarding CH₄-producing status, several studies suggested that factors such as age [23, 24], ethnic background [25, 26], gender [24, 27], exercise status [28], and various gastrointestinal diseases [29-34] play a role in the increased level of breath CH₄. Polag and Keppler [35] gave an overview of the variability of study parameters and calculated that 38% of humans globally are CH₄ producers, with an average CH₄ content of around 17 ppmv in breath air. For further discussion regarding microbial formation of CH₄ in the human gut system and the physiological factors that might control these processes, see the review articles by de Lacy Castello et al. [36] and Levitt et al. [20].

Hereafter, based on the stable carbon isotope and CH₄ emission patterns of various age groups, it was hypothesized that next to microbial sources in the gastrointestinal tracts there might be other, yet unidentified, endogenous cellular processes involved in CH₄ formation [18, 22, 37]. Indeed, long-term monitoring of CH₄ in combination with the observation of physical conditions suggested a relationship between deviations in breath CH₄ production from base level and immune reactions and inflammatory processes [37, 38]. Thus, there was increasing evidence that CH₄ has a bioactive role in the cellular physiology of eukaryotes and could be considered as a diagnostic marker of oxidoreductive stress [18, 39].

Alternative mechanism(s) of non-microbial methane formation in eukaryotes

In plants, it has been shown that CH₄ production is influenced by environmental factors such as visible light [4], ultraviolet-B radiation [5, 6, 40-44], and temperature [2, 6, 42, 45]. Several precursors of non-microbial CH₄ production, including methoxy groups of plant pectins [2, 5, 46], lignin, cellulose [45], ascorbic acid [47], L-methionine [8, 48], and epicuticular wax [49], have been suggested. In addition, for higher plants and microalgae, it has been demonstrated that environmental stressors enhance CH₄ formation drastically [14, 50].

To understand non-microbial CH₄ formation in eukaryotes and other organisms, it is important to consider the role of reactive oxygen species (ROS, e.g., hydroxyl radicals ($\cdot\text{OH}$), superoxide radicals (O^{2-}), hydrogen peroxide (H_2O_2), or carbonate radicals ($\text{CO}_3^{\cdot-}$)), iron species, and carbon precursor compounds. In highly oxidative environments generated *in vitro* by a chemical model system containing iron (II/III), H_2O_2 , and the radical scavenger ascorbate, CH₄ is readily formed from organosulfur and nitrogen compounds, with the highest conversion rates found for DMSO ([51]. Under these Fenton-type conditions, in the presence of H_2O_2 , nonheme oxo-iron(IV) ($[\text{Fe}^{\text{IV}}=\text{O}]^{2+}$) oxidizes methyl sulfides to sulfoxides, which then results in selective formation of methyl radicals by sulfoxide demethylation and ultimately leads to CH₄ [51, 52]. Alternatively, ROS can directly react with methyl sulfides to produce methyl radicals or peroxomethyl radicals in the presence of oxygen [53-55], subsequently resulting in CH₄ formation or oxidized C1-species such as methanol or formaldehyde.

The Fenton reaction, including the interaction of free iron species and H_2O_2 , is a key reaction in biological systems; its major cause is oxidative stress. In living cells, Fenton chemistry takes place as iron is an essential trace element [56] and H_2O_2 is produced during respiratory, and generally metabolic, activity [57]. Hydrogen peroxide and ferrous iron (Fe^{2+}) either react to ferric iron (Fe^{3+}), OH^- and $\cdot\text{OH}$ radicals, or, alternatively, to $[\text{Fe}^{\text{IV}}=\text{O}]^{2+}$ and water [58]. This provides the basis for our understanding of CH₄ formation in cells under oxic conditions. A wide spectrum of molecules that act as methyl donors for CH₄ formation are conceivable. However, of particular interest are compounds with sulfur (S) and nitrogen (N)-bonded methyl groups that arise during cellular metabolism or are externally provided. For methylated sulfur compounds, these include DMSO and methionine, which are ubiquitous in the environment [59]. In addition, methylated nitrogen compounds such as betaine, choline, or trimethylamine might also serve as CH₄ precursors. This has recently been proven for many organisms from the three domains of life using culture experiments, and a detailed reaction mechanism for CH₄ formation has been suggested, highlighting the interaction between ROS, iron, and S- and N-methylated compounds [17]. The authors also showed that oxidative stress led to increased CH₄ formation in the studied organisms.

In summary, the reaction of methylated sulfur compounds such as DMSO and methionine with Fenton type chemistry involving ROS, carbonate radicals, or oxo-iron(IV) results in the formation of methyl radicals ($\cdot\text{CH}_3$), of which a portion reacts to CH₄ through abstraction of a hydrogen atom from hydrocarbons, hydrogen peroxide, or hydrogen carbonate. Alternatively, the methyl radicals form oxidized C1-species such as methanol, formaldehyde, or formic acid. Thus, it is conceivable that there is *in vivo* formation of C1 compounds as a result of ROS formation and interaction with methylated compounds. Therefore, we consider the administration of isotopically labeled DMSO and methionine as ideal model compounds to confirm the occurrence of ROS-driven CH₄ formation in humans.

Application of DMSO to humans

Dimethyl sulfoxide (DMSO) is an organic polar aprotic molecule that was first synthesized in 1866. It was used as an important solvent for many decades before being proposed for use as a pharmaceutical in the 1960s by Stanley Jacob. Because of its ability to rapidly penetrate through human skin and its properties as a free-radical ($\cdot\text{OH}$) scavenger, it was widely used as an anti-inflammatory, antipain, and neuroprotective agent. A wide range of biological and pharmacological effects of DMSO is described by Jacob and Herschler [60] for the interested reader. Since 1978, DMSO has been approved by the United States Food and Drug Administration (FDA) for treatment of interstitial cystitis. Other medical applications, as well as potential physiological and pathological effects of DMSO, are highly controversially discussed. For example, Amemori et al. [61] found that oral administration of DMSO is an effective treatment for amyloid A amyloidosis. On the other hand, experiments with rats found that DMSO might induce retinal apoptosis [62]. Despite the differing results of the various studies, it is generally assumed that DMSO is nontoxic below 10% (v/v) [63] with an oral medium lethal dose of 28,300 mg/kg (rat) and a dermal medium lethal dose of 40,000 mg/kg (rat).

Aims and postulates

Recent results have shown that CH₄ might be formed in all organisms and that the formation of methyl radicals induced by ROS is a prerequisite for the generation of CH₄. The experiments described in this paper were undertaken in order to unambiguously demonstrate (as a first proof of principle) that CH₄ is endogenously formed in humans by a radical-driven process without involvement of the well-known microbial sources (methanogens) living in the gastrointestinal tract. Therefore, a volunteer—the first author of this study—applied isotopically labeled (²H or ¹³C) DMSO on the skin (arm), consumed it by mouth, and applied it to blood samples. In addition, the amino acid methionine (with an isotopically labeled ¹³C methyl group) was also applied to the blood samples. The released gases were analyzed for their isotopic composition to unambiguously identify formation of CH₄ from the precursor compounds DMSO and methionine. Based on the results and the formation patterns observed, we discuss several hypotheses concerning the origin of cell-based CH₄ production and its potential physiological role in mammals. Finally, as DMSO has already been investigated for many years while its potential positive role for medical use is highly uncertain, we will briefly discuss the potential application of DMSO to reveal and counteract oxidative stress.

2. Materials and Methods

2.1. Subject, materials, experiments, and sampling of air

2.1.1. Subject of the study

All experiments and measurements were conducted by the principal investigator (PI) and first author of this study (F.K.) from June 2018 to October 2020. The subject was a healthy 55-year-old man without known disease, prescribed medications, or drug intake. The average breath CH₄ production value of the subject was 9 ± 6.7 ppmv, measured within a period of 72 weeks [38], and he was thus classified as a medium to high emitter (see explanation above). Air and blood samples were provided by the PI as shown in section 2.1.3 below. A surrogate investigator (D.P.) was designated to obtain informed consent from the self-experimenter (F.K.), in agreement with the ethics relevant to solitary self-experimentation [64]. The work described has been carried out in accordance with The Code of Ethics of the World Medical Association. The research was reviewed by the Medical Research Council of Hungary (ETT-TUKEB) and it was approved as part of the protocol "Mapping metabolic pathways of endogenous gas formation by isotopic analysis of the gas composition of human samples" (6420-8-2023/EUIG/768).

2.1.2. Materials: Position-specific isotopically labeled DMSO and methionine

The isotopically labeled sulphur-bonded methyl group(s) in DMSO (¹³C₂-DMSO, 99%, Campro Scientific GmbH, Germany and DMSO-d₆, 99.9 atom %; Sigma Aldrich, Germany) and methionine (¹³CH₃-MET, Sigma-Aldrich, Taufkirchen, Germany; Isotec 99% ¹³C atoms) were investigated as methyl precursors for CH₄ (Fig. 1). Please note that ²H-labeled methionine was not available for experiments.

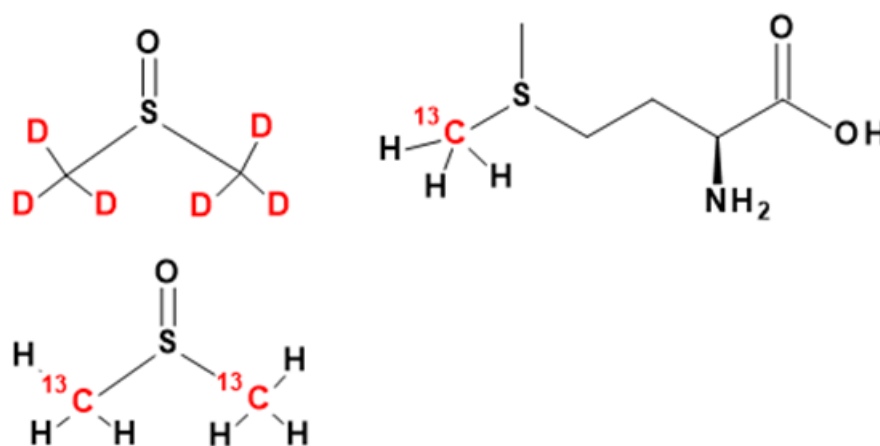


Figure 1. The chemical structures of DMSO (left) and methionine (right). The isotopically labeled carbon and hydrogen positions are indicated by ^{13}C and D (Deuterium, ^2H) highlighted in red font.

2.1.3. Experiments and sampling of air

A graphical representation of the set-up of the three individual experimental series (oral intake, arm exposure to sunlight, and blood experiments), including collection of samples and the applied measurements, is outlined in Figure 2. Table S1 shows the timeline of the experiments.

Oral intake of ^{13}C - and ^2H -labeled DMSO

The volunteer of the study swallowed 100 μL ^{13}C -CH₄ DMSO (4% ^{13}C -content, dissolved in 300 mL H₂O) or 1 mL of ^2H -CH₄ DMSO (10% ^2H -content, dissolved in 300 mL H₂O), respectively. Subsequently, breath CH₄ concentration and isotope values of CH₄ ($\delta^{13}\text{C}$ or $\delta^2\text{H}$, respectively) were monitored for a period of 130 minutes. The breath samples were collected using 1 L Tedlar bags. The breath CH₄ sampling procedure was performed in a consistent manner. During breath air collection, the volunteer breathed normally, stopped breathing for around 5 s, and then filled the Tedlar bag with expired air (range of 0.8 to 1 L). Depending on the study parameter, the gaseous sample was analyzed by CRDS, GC-FID, or GC-TC-IRMS immediately after sampling (see analytical measurements below).

Arm incubations and exposure to solar light

For CH₄ skin emission analysis, the forearm of the subject was placed inside a cylindrical chamber (see photo documentation 1 in Supplementary Material) made of PTFE foil (transparent for UV light) with a diameter of 18.5 cm and a length of 42.5 cm (volume = 11.7 L). The round opening at the back was sealed with an elastic PTFE foil tied to the chamber and fixed along the upper arm. A gas inlet and outlet PTFE tube system was attached to the chamber. The pressure of the chamber was constant during the whole monitoring phase. Ventilation at the inside front of the chamber provided homogeneous air mixture. The outlet tube was directly connected to the CRDS system (see analytical measurements below) for in situ online analysis of CH₄ and CO₂ concentrations and $\delta^{13}\text{C}$ values. First, the empty chamber (filled with laboratory air) was measured as a control. Next, the volunteer thoroughly washed his arm with tap water and dried it with a paper towel before placing it in the chamber for a period of 30 minutes to obtain a control value. Then, ^{13}C -CH₄-labeled DMSO (a mixture of 400 μl DMSO + 100 μl ^{13}C -labeled DMSO + 500 μl H₂O) was thoroughly distributed on the skin of the left upper forearm (penetrated area of around 30 cm²) and the air in the chamber was connected to the CRDS measurement system for a monitoring period of 1h. Afterwards, the forearm was exposed to natural solar light in the field for a period of 1h (from 10 to 11 am, in July in Heidelberg,

Germany). After returning from the field to the laboratory (within 5 minutes) the left arm was again placed in the chamber and monitored for changes in $\delta^{13}\text{C-CH}_4$ values for 1h. The same procedure was repeated the following two days and the untreated right arm served to record control values.

Blood samples and incubation with DMSO and methionine

Approximately 20 mL of whole blood samples were collected from the PI through venipuncture by using 4 x 7.5 mL S-Monovettes® containing EDTA to prevent coagulation. Samples were immediately processed for isotope label experiments. Therefore, ^{13}C -labeled DMSO and methionine were added to 1 mL of blood in autoclaved 40 mL headspace vials (Supelco 27184), so that the final concentration of the added compound was 1 mM or 10 mM. Vials were sealed using a hole-type screw cap (Supelco) fitted with a PTFE/silicone septum (Supelco). The control samples were prepared in the same way, except that the added DMSO and methionine were isotopically not enriched in ^{13}C . All samples were prepared in triplicates and incubated at 36°C for 24 hours before the gas phase in the vials was analyzed (first day). Afterwards, vials were opened to equilibrate with air in a fume cupboard. After 30 minutes, samples were again sealed with a PTFE/silicone septum and incubated at 36°C for 24 hours before the gas phase was analyzed (second day).

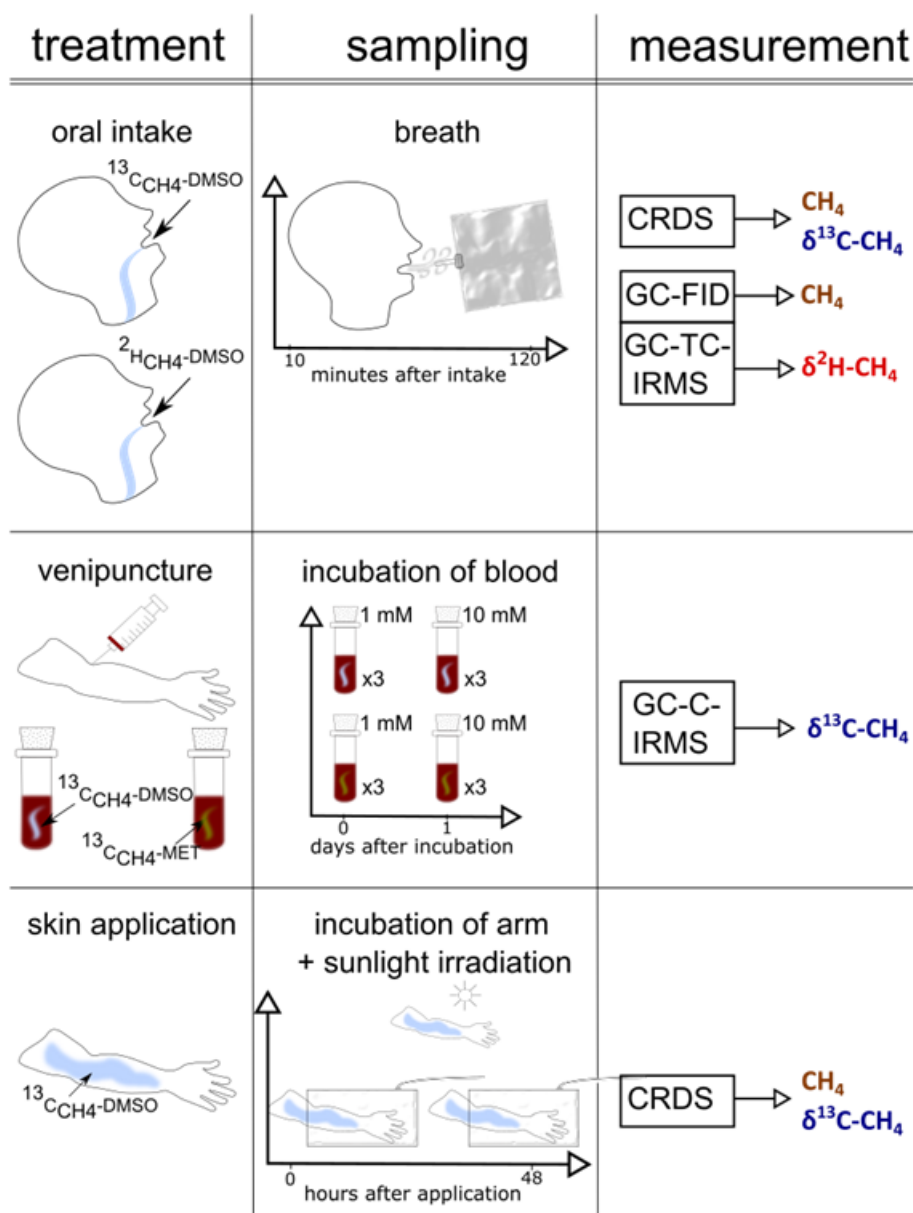


Figure 2. A graphical representation of the set-up (treatment, sampling, measurement) of the three experiments (skin, blood, and oral administration) performed in this study.

2.2. Analytical measurements

The analytical laser technique applied in this study to obtain online stable carbon isotope measurements and concentrations of CH₄ is almost identical to that described previously [22]. In addition, stable carbon and hydrogen isotope analysis were conducted by applying continuous flow isotope ratio mass spectrometry (GC-IRMS), as described in Einzmann et al. [65]. However, we will briefly describe the applied analytical techniques in the sections below. For more analytical details, and the application of stable isotope techniques, please refer to the studies by Keppler et al. [22] and Einzmann et al. [65], and to the Supplementary Material.

2.2.1. Natural abundance of ¹³C/¹²C and ²H/¹H, definition of δ values, isotopic excess, and Keeling method

Throughout this paper, the “delta” (δ) notation—the relative difference of the isotope ratio of a material to that of a standard V-PDB (Vienna Pee Dee Belemnite, ¹³C/¹²C ratio of 0.011108) or VSMOW (Vienna Standard Mean Ocean Water, ²H/¹H ratio of 0.00011576)—is used; values of δ¹³C and δ²H relative to those of V-PDB and VSMOW, respectively, are defined by the equation:

$$\delta^{13}\text{C} = \left(\frac{(^{13}\text{C}/^{12}\text{C})_{\text{sample}}}{(^{13}\text{C}/^{12}\text{C})_{\text{standard}}} - 1 \right) \quad (1)$$

$$\delta^2\text{H} = \left(\frac{(^2\text{H}/^1\text{H})_{\text{sample}}}{(^2\text{H}/^1\text{H})_{\text{standard}}} - 1 \right) \quad (2)$$

To comply with the guidelines of the International System of Units (SI), we followed the proposal of Brand and Coplen [66] and used the term *urey*, after H.C. Urey (symbol Ur), as the isotope delta value unit. In such a manner, an isotope-delta value expressed traditionally as -60‰ can be written as 60 mUr. For natural sources of CH₄, typical δ¹³C-CH₄ and δ²H-CH₄ values are in the range of -20 to -100 mUr [12] and -100 to -400 mUr [44], respectively.

The isotopic difference (Δ) between control and sample is defined as:

$$\Delta = \delta^{13}\text{C}_{\text{sample}} - \delta^{13}\text{C}_{\text{control}} \quad (3)$$

¹³C % and ²H % excess were calculated as follows:

$$\begin{aligned} {}^{13}\text{C} \% \text{ excess} &= \left(\left(\frac{{}^{13}\text{C}}{{}^{13}\text{C} + {}^{12}\text{C}} \right)_{\text{Labelled}} - \left(\frac{{}^{13}\text{C}}{{}^{13}\text{C} + {}^{12}\text{C}} \right)_{\text{Basis}} \right) * 100 \\ &= {}^{13}\text{C} \%_{\text{Labelled}} - {}^{13}\text{C} \%_{\text{Basis}} \end{aligned} \quad (4)$$

$$\begin{aligned} {}^2\text{H} \% \text{ excess} &= \left(\left(\frac{{}^2\text{H}}{{}^2\text{H} + {}^1\text{H}} \right)_{\text{Labelled}} - \left(\frac{{}^2\text{H}}{{}^2\text{H} + {}^1\text{H}} \right)_{\text{Basis}} \right) * 100 \\ &= {}^2\text{H} \%_{\text{Labelled}} - {}^2\text{H} \%_{\text{Basis}} \end{aligned}$$

2.2.2. Laser absorption spectroscopy - Cavity Ring Down Spectroscopy

Measurements of CH₄ concentrations and stable carbon isotope values

Cavity Ring Down Spectroscopy (CRDS) is a highly sensitive optical spectroscopic technique for measurements of both stable carbon isotopes values (δ¹³C-CH₄) and concentration of CH₄. The Tedlar-gas sample bag (from breath air) or the arm incubation Teflon chamber (see Supplementary Material) was connected to the CRDS, and flow rate to the analyzer was 23 mL/min. Before entering the analytical system, the gas was passed through two chemical traps filled with AscariteII® (sodium hydroxide coated silica) and Drierite® (anhydrous CaSO₄) to remove carbon dioxide (CO₂) and water, respectively. This was necessary due to the higher concentrations of CO₂ and water (up to 6%) in the breath sample that can cause inferences with the spectroscopic CH₄ measurements.

Stable carbon isotope values and concentrations of CH₄ were measured with a G2201-i cw-CRDS-Analyzer (Picarro, Inc., Santa Clara, USA). This instrument enables simultaneous measurements of CH₄ concentration, δ¹³C-CH₄ value, and water content in a gas sample. The concentration precision (1σ, 2 min average) specified by the manufacturer is 50 ppbv + 0.05% of reading (¹²C) and 10 ppbv + 0.05% of reading (¹³C) in the high dynamic range mode, and 5 ppbv + 0.05% of reading (¹²C) and 1 ppbv + 0.05% of reading (¹³C) in the high precision mode. The δ¹³C-CH₄ precision provided by the manufacturer is < 0.8 mUr. However, typical standard deviations (SD) of measurements of breath samples and standards (filled in Tedlar bags) were in the range of ± 1.2 ppbv and ± 0.3 mUr (1σ, 2 min average measurement interval) for concentration and stable isotope measurements, respectively (see also [22]).

In order to quality assure the δ¹³C-CH₄ values, some gas samples were measured both by CDRS and gas chromatography-combustion-isotope ratio mass spectrometer (GC-C-IRMS) (for details see section 2.2.4.1 below). Samples measured by IRMS were analyzed three times (n = 3) and the average standard deviations of the analytical measurements were in the range of 0.1 to 0.3 mUr. The measured difference between the two analytical systems was used to normalize the isotope data of the CRDS.

2.2.3. Measurements of CH₄ concentrations using gas chromatography flame ionization detection (GC-FID)

An aliquot (5 mL) of headspace gas was taken from the incubation vials (40 mL) or gas bags (1 L) using a gastight syringe. Before entering the analytical system, the gas sample was passed through a chemical trap filled with Drierite® to remove water. The sample gas was separated by gas chromatography using a GC-14B (Shimadzu, Japan) equipped with a 2 m column (Ø = 3.175 mm inner diameter), packed with Molecular Sieve 5A 60/80 mesh from Supelco. Methane was recorded by a flame ionization detector (FID), and its concentration was quantified by using two reference gases containing 9837 ppbv and 2192 ppbv CH₄.

2.2.4. Continuous flow isotope ratio mass spectrometry

2.2.4.1. Measurement of δ¹³C-CH₄ values

Gas from the Tedlar-gas bags (from breath samples) or 40 mL glass vials (headspace of blood samples) was transferred to an evacuated sample loop (40 mL). Interfering compounds were separated by GC and CH₄ trapped on Hayesep D. Afterwards, CH₄ was separated from interfering compounds by GC and transferred to a gas chromatography-combustion-isotope ratio mass spectrometer (GC-C-IRMS, Deltaplus XL mass spectrometer, ThermoQuest Finnigan, Bremen, Germany) via an open split. The working reference gas was CO₂ of high purity (carbon dioxide 4.5, Messer Griesheim, Frankfurt, Germany) with a known δ¹³C value of -23.64 mUr (calibrated at MPI for Biogeochemistry in Jena, Germany). All δ¹³C-CH₄ values were corrected using two CH₄ working standards (isometric instruments, Victoria, Canada) calibrated against NIST and IAEA reference substances. The calibrated δ¹³C-CH₄ values of the two working standards in mUr vs. V-PDB were -23.9 ± 0.2 and -54.5 ± 0.2. All samples were normalized by two-scale anchor calibration according to Paul et al. (2007).

2.2.4.2. Measurement of δ²H-CH₄ values

δ²H-CH₄ values were determined via GC-TC-IRMS. The same analytical set-up was applied as for stable carbon isotope measurements (see section 2.2.2 above) with the following modifications: The flow rate was 0.6 mL min⁻¹ and instead of combustion to CO₂ and H₂O, CH₄ was thermolytically converted (at 1450°C) to produce hydrogen (H₂) and carbon. After IRMS measurements of the hydrogen, the obtained δ²H values were normalized using two reference standards of high-purity CH₄ with δ²H values of -190.6 ± 0.2‰ (in-house) and -149.9 ± 0.2‰ (T-iso2, Isometric Instruments).

2.3. Statistics

Data analysis was performed using R 4.1.2 software. For data smoothing, the Loess method was used. For data analysis with CRDS (sampling rate = 1 s), mean values were taken for those periods in which data variation was less than 5% (measurement periods of 20–30 minutes). Calculation of $\Delta\delta^{13}\text{C}-\text{CH}_4$ values for both experiments are presented as the arithmetic means of the respective replicates together with their standard deviation (SD). Arithmetic means and SDs were calculated using Microsoft Excel (Microsoft Excel for Office 365 MSO).

3. Results

3.1. Oral intake of isotopically labeled DMSO and measurements of breath air

Figure 3 shows the breath CH_4 production and isotope difference of $\delta^{13}\text{C}-\text{CH}_4$ values (Fig. 3a) and $\delta^2\text{H}-\text{CH}_4$ values (Fig. 3b), relative to control values, monitored by breath air of the subject after oral intake of isotopically labeled (with ^{13}C and ^2H , respectively; see methods) DMSO over a time period of 130 minutes. Breath CH_4 production (with laboratory background values subtracted) during the monitoring period showed mean values of 2.2 ± 0.06 ppmv (Fig. 3a, top) and 12.4 ± 2.3 ppmv (Fig. 3b, top), respectively.

In the beginning and at the end of the experiment (after 130 minutes) the $\Delta\delta^{13}\text{C}-\text{CH}_4$ and $\Delta\delta^2\text{H}-\text{CH}_4$ values of around 0 mUr closely reflected the average isotopic signature of the volunteer's breath $\delta^{13}\text{C}-\text{CH}_4$ values and $\delta^2\text{H}-\text{CH}_4$ values. For details regarding the determination of the volunteer's isotope CH_4 source signatures observed for several periods in 2018 and 2019 without any treatment of isotopically labeled compounds, see the Supplementary Material. Within a few minutes of oral intake of isotopic labeled DMSO, the $\Delta\delta^{13}\text{C}-\text{CH}_4$ and $\Delta\delta^2\text{H}-\text{CH}_4$ values substantially increased towards less negative respective $\delta^{13}\text{C}/^2\text{H}-\text{CH}_4$ values, resulting in positive Δ values and becoming clearly distinguishable from background values. A maximum increase in Δ values of around 2.5 mUr and 4000 mUr for carbon and hydrogen, respectively, was found between 60 and 40 minutes. Afterwards, $\Delta\delta^{13}\text{C}-\text{CH}_4$ and $\Delta\delta^2\text{H}-\text{CH}_4$ values steadily decreased, almost reaching the initial $\delta^{13}\text{C}/^2\text{H}-\text{CH}_4$ values after 130 minutes.

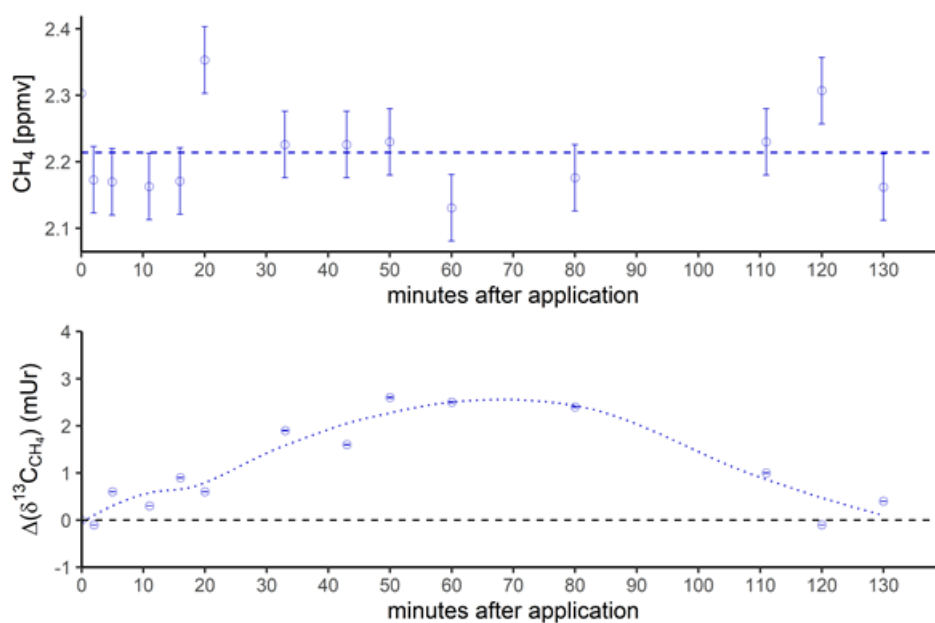


Figure 3. a. Breath CH_4 production (top) and isotope difference of $\delta^{13}\text{C}-\text{CH}_4$ values relative to control (bottom) after oral intake of isotopically labeled ^{13}C DMSO. The dashed line in the upper figure represents the CH_4 mean value. Error bars represent a sum of analytical uncertainties and statistical errors from triplicate measurements.

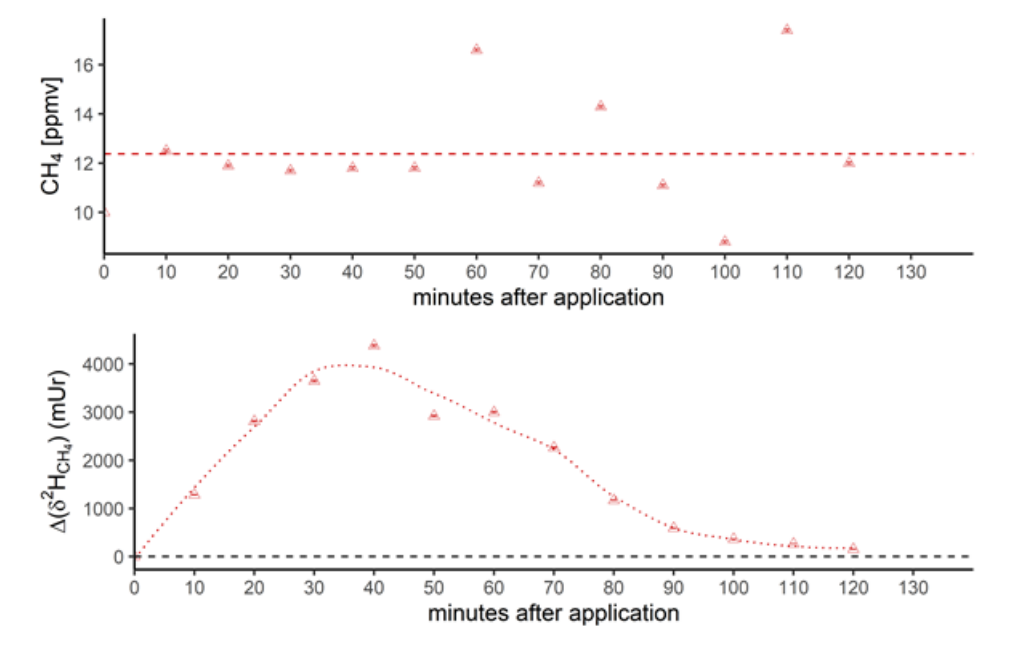


Figure 3. b. Breath CH₄ production (top) and isotope difference of δ²H-CH₄ values relative to control (bottom) after oral intake of isotopically labeled ²H DMSO. The dashed line in the upper figure represents the CH₄ mean value. Error bars represent a sum of analytical uncertainties and statistical errors from triplicate measurements, and lie within symbols.

In addition, Figure 3c compares the excess of isotopic label in released breath CH₄ from the supplemented ¹³C- and ²H-DMSO. The excess in both ¹³C-CH₄ and ²H-CH₄ gradually increased, with maximum values observed at 40 minutes for ²H-CH₄ (~0.68‰) and 50 minutes for ¹³C-CH₄ (~0.028‰). Both isotope tracers evidently indicate partial conversion of the methyl group of DMSO to CH₄ by processes within the human body. The maximum excess found for ²H-CH₄ was by a factor of around 24 higher when compared with ¹³C-CH₄. In this context, it should be noted that the amount of applied isotope ¹³C-labeling of DMSO was much lower for ¹³C experiments (see discussion).

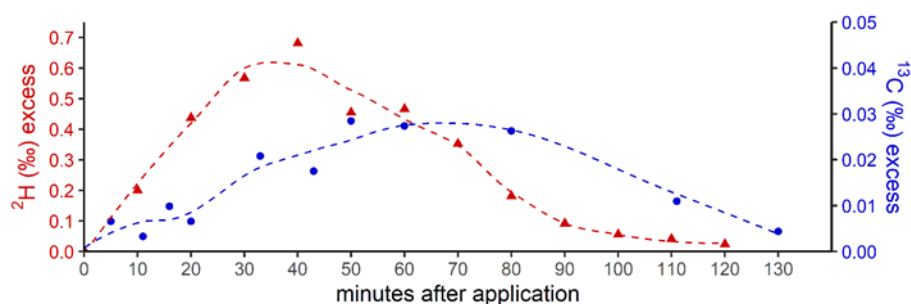


Figure 3. c. ²H (red) and ¹³C (blue) isotopic excess after the oral administration of labeled ¹³C and ²H DMSO, respectively.

3.2. Blood samples and addition of isotopically labeled DMSO and methionine

The supplementation of ¹³C-labeled DMSO and methionine at equimolar concentrations of 1 mM to the blood samples incubated for 24 hours (first day) at 36°C resulted in mean Δδ¹³C-CH₄ values of 95 ± 36 mUr and 2.2 ± 0.5 mUr for DMSO and methionine, respectively (Fig. 4). Repeated measurements of the same samples (after equilibration with laboratory air, see methods) and another incubation period of 24 hours (second day) exhibited lower mean Δδ¹³C-CH₄ values, ranging from 70

± 10 mUr and 0.24 ± 0.4 mUr for DMSO and methionine, respectively. Application of ten-fold higher concentrations of DMSO and methionine (10 mM) enhanced the formation of isotopically labeled CH_4 , with $\Delta\delta^{13}\text{C}-\text{CH}_4$ values ranging from 748 ± 362 mUr and 4.9 ± 3.5 mUr for DMSO and methionine, respectively. Again, repeated measurements of the same samples after another incubation period of 24 hours (second day) exhibited lower mean $\Delta\delta^{13}\text{C}-\text{CH}_4$ values, ranging from 588 ± 10 mUr and 1.4 ± 0.1 mUr for DMSO and methionine, respectively. Thus, the change in 10-fold concentrations was closely reflected by the change in $\Delta\delta^{13}\text{C}-\text{CH}_4$ values (factor of ~ 8) for both days, whilst for methionine the change in $\Delta\delta^{13}\text{C}-\text{CH}_4$ values was lower (factor 2.2 and 5.7 for day 1 and day 2, respectively). All control samples including blood without addition of isotopically labeled compounds did not show any measurable difference in $\delta^{13}\text{C}-\text{CH}_4$ values over the incubation time.

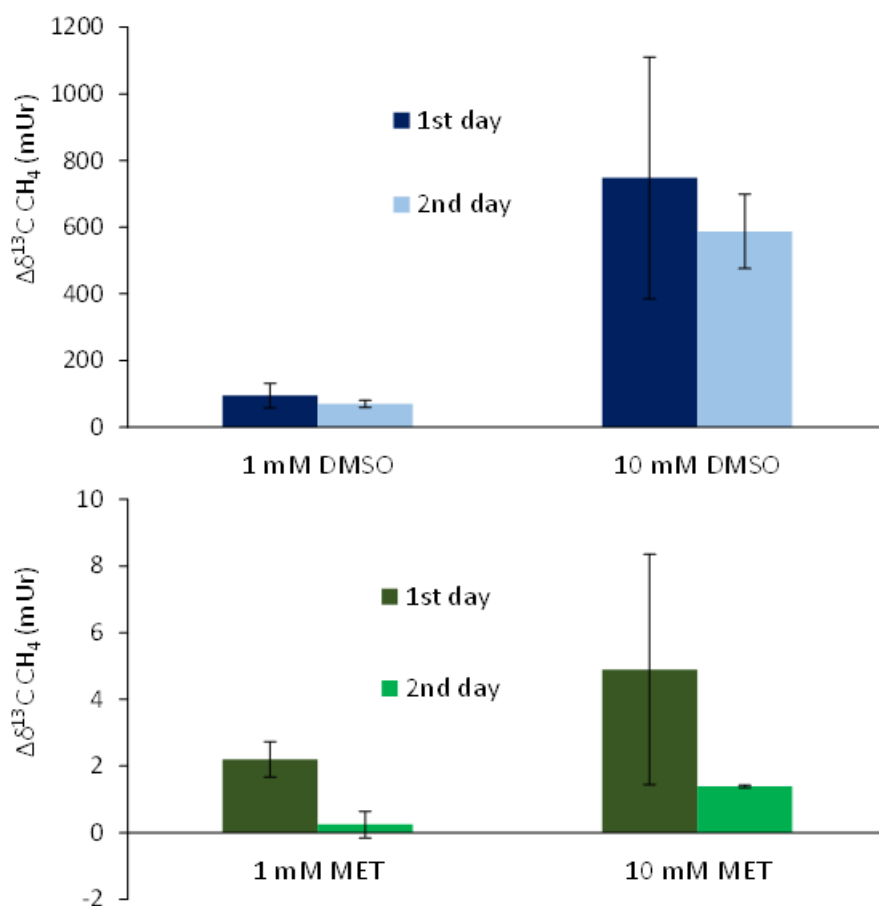


Figure 4. Isotope difference (relative to control values) of $\delta^{13}\text{C}-\text{CH}_4$ values of headspace-air after treatment of blood samples with ^{13}C -labeled DMSO (a) and methionine (b) incubated over two days at temperature of 36°C . Mean values of three replicated experiments ($n = 3$) are shown and error bars mark the SD. Controls samples (blood without addition of isotopically labeled compounds) did not show any measurable difference in $\delta^{13}\text{C}-\text{CH}_4$ values over the incubation time and thus are not graphically illustrated.

3.3. Skin application of isotopically labeled DMSO and incubation of arm including exposure to natural sunlight

Figure 5 shows the isotope difference of $\delta^{13}\text{C}-\text{CH}_4$ values relative to control values after the application of ^{13}C -labeled DMSO on the left forearm. Subsequent to DMSO application, $\delta^{13}\text{C}-\text{CH}_4$ -values increased by 30 mUr within 1 hour. After the volunteer exposed his left forearm to natural sunlight in the field, a maximum $\Delta\delta^{13}\text{C}-\text{CH}_4$ value of 50 mUr was observed. Please note that direct measurements during exposure to sunlight in the field were not possible. For experimental details,

we refer the reader to the method section. After 24 hours, $\delta^{13}\text{C}\text{-CH}_4$ -values measured for CH_4 release from skin of the left forearm were still enriched by 4 mUr whilst control values (incubation of the untreated right forearm) did not show any measurable changes. Again, exposure to sunlight in the field and subsequent laboratory measurements of CH_4 release from the skin of the forearm increased the $\Delta\delta^{13}\text{C}\text{-CH}_4$ value to 6 mUr. After 48 hours, $\delta^{13}\text{C}\text{-CH}_4$ values monitored from the release of the skin still showed a marginal but measurable ^{13}C enrichment of 1 mUr. After sunlight irradiation, no measurable increase of $\delta^{13}\text{C}\text{-CH}_4$ values was noted. The associated CH_4 concentrations of the chamber measurement series showed changes in the range of 1.96 to 2.08 ppmv which were close to the variations observed for the control measurements.

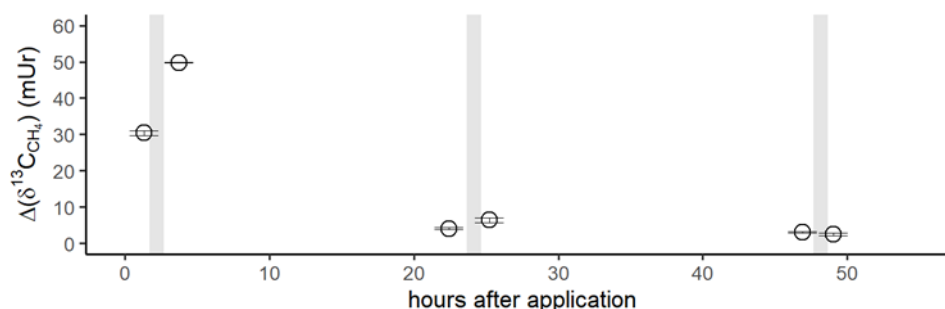


Figure 5. Isotope difference of $\delta^{13}\text{C}\text{-CH}_4$ values ($\Delta\delta^{13}\text{C}\text{-CH}_4$) relative to control values after the application of ^{13}C DMSO on the left forearm of the volunteer. Grey vertical lines represent periods of exposure to sunlight.

4. Discussion

Conversion of methylated sulfur compounds to methane

The three sets of experiments—involving the application of two potential CH_4 precursor compounds, DMSO and methionine, with isotopic labels—provide independent lines of evidence for partial conversion of the supplemented methyl group to CH_4 in the human body. The combination of the three experiments (oral intake, blood incubations, and skin application) were undertaken to confirm that CH_4 is endogenously formed in humans by a ROS-driven process without involvement of the well-known microbial sources (methanogens) occurring under anoxic conditions in the gastrointestinal tracts. However, we are aware that it is almost impossible to exclude the contribution of microbes during the screening of humans for CH_4 emissions.

Oral administration of ^{13}C -labeled DMSO

The measured isotopic changes for the two labeling experiments (Fig. 3) unambiguously demonstrated that the methyl group of DMSO is converted to CH_4 . The ^2H - and ^{13}C -excess values indicated that only a marginal fraction (0.68‰ and 0.028‰) of the CH_4 concentration measured in the subject's breath air (~2 to 16 ppmv) was actually derived from the isotopic labeled precursor methyl groups of DMSO. The observed variabilities in concentrations during the individual experiments (Fig. 3a and 3b, top panels) were in the range of intraday fluctuations. The observed difference in CH_4 base levels of approximately 10 ppmv between the experiments with ^{13}C DMSO and ^2H DMSO reflected usual changes in the individual's breath CH_4 state, as the two experiments were performed a few months apart. For details regarding variabilities of CH_4 base levels of the volunteer, see Polag and Keppler [37, 38]. The small concentration changes indicated by the supplementation of ^{13}C -labeled DMSO would be nondetectable when using conventional measurement techniques, and can only be traced using isotopic labeling techniques. To better compare the conversion of the two labeling approaches, it is necessary to consider the $^2\text{H}/^{13}\text{C}$ excess values, as shown in Figure 3c. The calculated maximum excess values of $^2\text{H}\text{-CH}_4$ and $^{13}\text{C}\text{-CH}_4$ (occurring at around 40 to 50 minutes) were ~0.68‰ and 0.028‰, respectively, and thus the maximum excess found for $^2\text{H}\text{-CH}_4$ was by a factor of around 24 higher when compared with ^{13}C -

CH₄. Please note that the ²H-CH₄ excess value of 0.68‰ includes three deuterium atoms from a ²H-labeled methyl group and a fourth, unlabeled hydrogen atom (see Fig. 6). To correct for this effect the maximum excess of ²H increases to 0.91‰ and the differences between excess values of ²H-CH₄ and ¹³C-CH₄ changes to a factor of 32. This value closely reflects the relationship of orally administered ²H and ¹³C isotope tracers (factor of 34). The reason for applying different amounts of ²H/¹³C DMSO isotopic labels was due to financial issues, as ²H-labeled DMSO is considerably cheaper than ¹³C-labeled DMSO. Nevertheless, both isotope tracers independently and clearly indicated similar conversion rates of the methyl group of DMSO when normalized to the amount of applied isotopic tracer. We suggest that the observed CH₄ formation is indicative of the formation of methyl radicals from DMSO induced by hydroxyl radicals or oxo-iron(IV) species, as recently proposed by Ernst et al. [17], Benzing et al. [52], and Althoff et al. [51] for biological and abiotic chemical systems. Once methyl radicals are formed, they can react with a hydrogen atom from hydrocarbons, hydrogen peroxide, or hydrogen carbonate to form CH₄. Formation of ¹³C-enriched CH₄ was already measurable a few minutes after the oral intake of the labeled substance for both isotope labeling experiments (²H and ¹³C). However, around 2 hours after oral administration, CH₄ formation from DMSO was barely detectable in breath air, potentially implying that most of the DMSO is converted in the human body within this timespan. A possible decay mechanism is the conversion of DMSO to dimethyl sulfide (DMS) by the molybdoenzyme DMSO reductase, which is widespread in all domains of life [67]. A currently proposed mechanism of DMSO reductase can be found in Le et al. [68].

Supplementation of ¹³C-labeled DMSO and methionine to blood samples

The experiments with blood samples were conducted to further demonstrate non-microbial formation of CH₄ when different S-methylated compounds were supplemented. When equimolar amounts of DMSO and methionine were added to the blood samples, the conversion of S-methyl bonded groups to CH₄ was much higher for DMSO than for methionine, with factors ranging from 43 to 423. It is well known that DMSO is a potent hydroxyl radical scavenger [69] forming CH₄, ethane, and oxidized C1 compounds such as formaldehyde and formate, depending on the experimental conditions [70-72]. The observed differences between application of DMSO and methionine are in line with previous experiments conducted by Althoff et al. [51] and Ernst et al. [17], who showed the preferential formation of ROS-induced formation of CH₄ from DMSO relative to methionine in chemical systems and living organisms, respectively. However, in our study the difference between DMSO and methionine was even more pronounced, and might be explained by the specific composition of the blood samples, i.e., amounts and availability of iron species and ROS. In addition, methionine needs to be oxidized to methionine sulfoxide before the methyl groups can be cleaved off [51]. Human blood and plasma contain high amounts of iron species, particularly in the form of hemoglobin, and the range of H₂O₂ might be in a normal concentration range of 1–5 μM but increase 30–50 μM during chronic inflammation in certain disease states [73]. Thus, the interplay between iron species and ROS in blood might be highly supportive for the formation of CH₄, given that the required methyl precursor compounds are also available. Interestingly, ten-fold higher DMSO supplementation was well reflected by the amounts of formed labeled CH₄ (factor of ~8), whilst a considerably lower increase was observed (mean factor of ~4) for the addition of methionine. It is also obvious that CH₄ formation from DMSO was observable for much longer (at least for 48 hours) in the blood samples when compared with oral administration of DMSO (see section above), indicating that different degradation processes in the human body might have contributed to the observed pattern.

Dermal CH₄ emissions after treatment of isotopically labeled DMSO

The application of ¹³C-labeled DMSO on the volunteer's forearm clearly showed release of isotopically labeled CH₄ immediately after incubation of the penetrated skin section (Fig. 5) under laboratory conditions. Based on our current understanding—including knowledge on ROS-driven CH₄ formation, and that DMSO rapidly penetrates through human skin—this observation is highly

indicative of methyl radical formation induced by ROS that occurs in the epidermis or dermis of the skin. There is frequent formation of ROS in the cells and it is well known that skin exposure to light—including wavelengths of visible light, UVA/UVB, and IR—induces and increases ROS levels [74, 75]. After the volunteer exposed his left forearm to natural sunlight in the field for 1 hour, a strong isotope change in $\delta^{13}\text{C}\text{-CH}_4$ -values (~70% higher relative to laboratory light exposure) was measured, even though these measurements were conducted after the exposure of direct solar radiation. This implies that enhanced levels of ROS were caused by the irradiation of solar light, leading to the formation of CH_4 , which could only be made visible by the administration of ^{13}C -labeled DMSO. After around 24 hours, the release of ^{13}C -labeled CH_4 from the skin under laboratory incubation conditions was still measurable, and increased again (by about 50%) after the exposure of the skin to natural sunlight. When repeating the same procedure after 48 hours, a small but indicative change in $\delta^{13}\text{C}\text{-CH}_4$ -values was still observed for the laboratory exposure incubations of the forearm. No additional increase of $\delta^{13}\text{C}\text{-CH}_4$ -values could be measured for the effect of natural sunlight. However, it was remarkable to observe DMSO-related liberation of CH_4 from the skin even 50 hours after application of ^{13}C -labeled DMSO. There exist only a few studies that deal in detail with the release of CH_4 from human skin, and in general these emissions are considered to be much smaller than those measured for breath release [35]. This was recently confirmed by Li et al. [76], who quantified dermal and exhaled CH_4 of 20 volunteers using climate chambers and reported that the average estimated exhaled CH_4 release rate was about 19 (max. range 13–37) times higher than the average dermal CH_4 emission rate. For completeness it should be noted that Mochalski et al. [77] measured emission rates of selected volatile organic compounds from the skin of healthy volunteers. However, the researchers did not detect CH_4 as they screened for larger carbon compounds, including C4 to C10 substances, and found relatively large emissions for three volatiles: acetone, acetaldehyde and 6-methyl-5-hepten-2-one.

ROS-induced non-microbial formation of CH_4 from methylated S-/N-compounds in humans: A hypothesis

The observed formation of CH_4 from the S-bonded methyl groups of DMSO or methionine provides strong support for a radical-driven process of CH_4 formation. Based on the three applied isotopic labeling experiments and a previous study demonstrating ROS-driven CH_4 formation from in vitro experiments of many organisms [17], we propose a reaction scheme showing the interplay of methyl precursors, ROS, and iron species that eventually leads to formation of CH_4 in humans (Fig. 6).

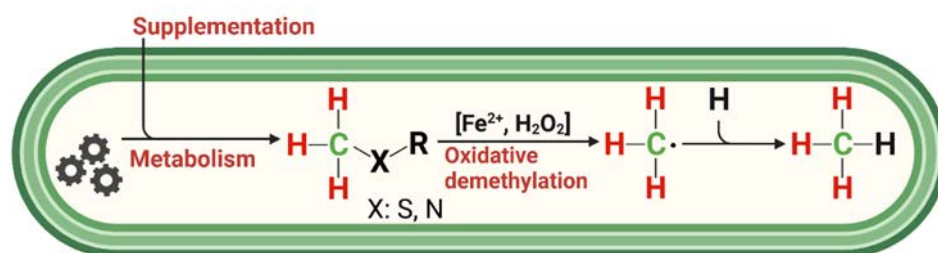


Figure 6. Simplified reaction scheme for endogenous CH_4 formation in humans. Methylated S-/N-compounds produced by metabolism or externally supplemented act as $\cdot\text{OH}$ scavengers or react with oxo-iron(IV) ($[\text{Fe}^{\text{IV}}=\text{O}]^{2+}$) to produce methyl radicals. Activation of hydrogen peroxide by ferrous iron (Fenton systems) leads to several oxidizing agents, such as $[\text{Fe}^{\text{IV}}=\text{O}]^{2+}$ or hydroxyl radicals, depending on the reaction conditions. Subsequently, CH_4 is formed through the reaction of a methyl radical with a hydrogen atom derived from hydrocarbons, hydrogen peroxide, or hydrogen carbonate. Red color and green color indicate hydrogen and carbon atoms, respectively, of methylated sulfur compounds labeled with ^2H and ^{13}C , as applied in this study to subsequently trace the formation of CH_4 in humans.

The three major players in this reaction scheme are ROS, iron, and methyl groups bonded to sulfur and nitrogen compounds. Below, we will briefly summarize their role in humans with respect to non-microbial CH₄ formation.

Initially considered as principally toxic, today ROS are well-known for having beneficial or deleterious effects in aerobic organisms [57, 78-80]. The concentration of H₂O₂ in the normal cytoplasm, mitochondrial matrix, and endoplasmic reticulum (ER) lumen varies by several orders of magnitudes (from 80 pM to 700 nM) [81] and is even higher in blood and plasma at normal concentrations of 1–5 μM, but increases to 30–50 μM during chronic inflammation in certain disease states [73]. On the one hand, ROS play various roles in the cellular functioning of aerobic organisms, are involved in many redox-governing activities of the cells for the preservation of cellular homeostasis, and are required for many important signaling reactions. On the other hand, elevated ROS levels can lead to severe damage in cells. In this context, it has been suggested that frequently increased oxidative stress leads to an overproduction of ROS, causing many diseases and a variety of age-related disorders such as Parkinson's disease, Alzheimer's dementia, chronic inflammatory diseases, atherosclerosis, heart attacks, cancer, ischaemia/reperfusion injury and arteriosclerosis. Thus, it can be easily envisaged that CH₄ might be formed at highly fluctuating levels in different organelles and might potentially serve to monitor enhanced ROS levels in humans. This hypothesis is supported by several recent monitoring studies:

(I) The observation that breath CH₄ levels increase with advanced age [24] might be an indication of the human age-related increase of systemic inflammation accompanied by enhanced ROS levels. (II) Long-term monitoring studies of breath CH₄ from several volunteers provided evidence that abrupt deviations in breath CH₄ levels from baseline were linked to inflammatory processes and immune reactions [37]. In this context, infectious diseases were mostly accompanied by temporary elevated breath CH₄ formation. Next, it was hypothesized that vaccinations as induced perturbations of the immune system might cause substantial fluctuations in the breath CH₄ level of people, indicating individual immune responses and immune states. (III) This has recently been proven by Polag and Keppler [38], who investigated the breath CH₄ levels of 12 volunteers after Covid-19 vaccination. They clearly found large deviations from the average breath CH₄ values of the subjects after vaccination and concluded that these deviations were likely related to immune reactions and may also have originated from redox homeostasis in cells. A change in the breath CH₄ levels from individual baseline values could be used to monitor changes in levels of ROS and oxidative stress, and could potentially be used to classify immune responses. (IV) Finally, Tuboly et al. [82] investigated the possibility of CH₄ generation in low-CH₄ emitters that consumed high doses of ethanol with the aim to increase oxidative stress. A transient, significant CH₄ production was noted after excessive ethanol intake. The researchers found similar results when they repeated the ethanol experiments with rats. They further investigated the hypothesis that L-alpha-glycerolphosphorylcholine (GPC) may influence CH₄ formation through the modulation of alcohol-induced mitochondrial dysfunction.

This brings us to the next point: to counteract oxidative stress, aerobic cells possess many antioxidative systems that function to keep the ROS level in a non-toxic range. Methyl precursors—particularly those where the methyl group is bonded to sulfur and nitrogen compounds—can readily be cleaved off to produce CH₄ or oxidized C1-species [17, 51, 52]. The various available S-/N-methylated compounds in biological systems will cause different efficiency of CH₄ production and consumption of ROS. DMSO is not produced in humans and is only consumed via the diet in relatively small quantities [83]. However, this effective radical scavenger was ideally suited to test the hypothesis of non-microbial CH₄ formation in humans. It is non-toxic in the applied doses, penetrates rapidly through human skin, and is easily distributed in the body, as it dissolves in both polar and nonpolar compounds. On the other hand, the other applied S-methylated compound, methionine, is an essential amino acid in humans that has an important role in metabolism and health. It is the precursor of other important compounds, such as cysteine, S-adenosyl methionine (SAM), and glutathione. It has also been shown to produce CH₄, albeit at much lower conversion rates when compared with DMSO.

Nitrogen-methylated substances such as choline (2-Hydroxyethyl-trimethylammonium) are formed in humans but are also essential compounds for maintaining health. Therefore, they must be consumed by diet as choline or as choline phospholipids. Large amounts of choline are stored in the human cell membranes and organelles as phospholipids, and inside cells as phosphatidylcholines and GPC. Choline has been shown to form CH₄ in a chemical model system containing iron and hydrogen peroxide [51, 84] but this was not confirmed in bacterial culture experiments [17]. Tuboly et al. [82] showed that exogenous GPC protected against ethanol-induced mitochondrial electron transport chain dysfunction in rat liver, the primary target of alcohol-induced oxido-reductive stress. Therefore, exogenous addition of methylated compounds might strongly increase CH₄ production and ROS consumption. In this context, it is of interest to further discuss the potential role of DMSO as an effective scavenger of radicals to counteract enhanced oxidative stress induced by ROS. DMSO has already been investigated for many years, but its beneficial role for medical use remained highly uncertain (see introduction).

Finally the concentrations of free iron (in the form of iron(II)) is of importance for enhanced production of hydroxyl radicals (Fenton type reactions) in biological systems [57]. However, inappropriately low or high levels of iron are detrimental and contribute to a wide range of diseases [85], so understanding the dysregulation of iron metabolism is crucial in the search for therapeutics [86]. Harmful oxidative distress could be observed in states of both iron deficiency (anemia) and overload (ferroptosis) [87]. It is plausible that appropriate supplementation of iron is beneficial to health, which may be related to its role in contributing to the homeostasis of cellular ROS through production of CH₄.

A detailed understanding of the interplay between ROS, iron, and methylated substrates in humans is necessary to better understand radical-driven CH₄ and to answer the question if cellular formation of CH₄ has a physiological role in humans. In this context, monitoring CH₄ as an indicator for ROS-driven processes could be a promising approach in biochemical research, and breath CH₄ could be used as a diagnostic tool in the fields of system biology and precision medicine. This could include the application of isotopic labeling experiments of methylated precursor substances (with a ¹³C or ²H label), as this approach could specifically visualize ROS-related CH₄ generation and thus overcome the problem of higher breath CH₄ background concentrations derived from microbial sources. These changes may be interesting for diagnostic purposes. Moreover, the possibility exists that such changes may affect the overall cellular response to intracellular hypoxia. Simple asphyxiants, such as CH₄, act by physically limiting the utilization of oxygen and can modify the symbiosis with other gaseous compounds within the internal milieu of aerobic cells. Although CH₄ is conventionally believed to be physiologically inert, a comprehensive view of its biological effects in various hypoxic and inflammatory scenarios has been demonstrated [88]. Notably, it has been shown that CH₄ can modulate the pathways involved in key events of inflammation via master switches, such as Nrf2/Keap1 and NF-κB (for a review, see [18]). Several studies have also demonstrated that exogenous CH₄ modulates the intrinsic, mitochondrial pathway of pro-apoptotic activation in model experiments [89]. Furthermore, sequential in vitro studies with exogenous normoxic CH₄ in simulated ischemia-reperfusion environments provided evidence that CH₄ preserves the mitochondrial respiratory capacity in cells exposed to anoxia [90]. In a similar protocol, CH₄ treatment restricted the forward electron transfer within Complex I in control mitochondria while effectively restricting reverse electron transport (RET) in post-anoxic mitochondria. In parallel studies, CH₄ influenced several components of the endoplasmic reticulum-mitochondria-related pro-apoptotic signaling pathways, the oxidative phosphorylation capacity was more preserved, and the relative mRNA expression for hypoxia- and ER stress-associated genes (including HIF-1α) was significantly reduced [91]. For detailed discussion regarding the potential applications of monitoring CH₄ in medical research and health sciences, see [18, 38, 92, 93].

5. Conclusions

We are aware that the investigation of ROS-driven cellular CH₄ formation in only one subject is too low to draw broad and general conclusions. However, this study represents a first proof of

concept that cellular CH₄ formation occurs in the human body and is most likely a result of the interplay between ROS and methylated substrates. This process can currently only be made clearly visible by applying stable isotope tracing techniques to distinguish CH₄ formation in humans from that of microbes living in the gastrointestinal tract. Together with other recently published studies [17, 38, 92] it is becoming obvious that ROS-driven CH₄ formation might be a necessary phenomenon of aerobic life. Consequently, non-microbial aerobic CH₄ formation should be highly variable in time and source strength, as it may be an integral part of the cellular responses towards changes in oxidative status present in humans. Large changes in human breath levels have been observed by several recent monitoring studies [38, 76, 82], and some have suggested that variations in CH₄ breath levels are unlikely to be explained by microbial formation in the human digestive system. However, additional investigations are required to obtain unambiguous evidence of non-microbial CH₄ formation in humans and the underlying processes of its generation. This will be a significant challenge, because in the case of high emitters— where CH₄ formation by methanogens is the dominant process—it is difficult to distinguish between the non-microbial and microbial pathways of CH₄ production. Nevertheless, for low and medium CH₄ emitters, who comprise about 70% of the global population, we suggest that substantially changed human breath CH₄ levels from individual baseline values may be used to detect changes in oxidative stress and ROS levels, and could potentially be used to classify immune responses, as recently suggested by Polag and Keppler [38]. Therefore, future investigations should focus on deciphering the potential physiological role of CH₄ formation in humans, as well as on the monitoring of CH₄ as indicator for individual immune states and a potential biomarker of oxidative stress. In addition, revisiting and studying in greater detail the potential role of DMSO as an effective ·OH scavenger, and its use for human medical research, might be worthwhile.

Supplementary Materials: The Supplementary Material for this article can be downloaded at XXX.

Author Contributions:

F.K. conceived the study, conducted the experiments, and analyzed the samples. F.K. and D.P. discussed and evaluated the data and prepared the graphical illustration of the data. M.B. provided critical discussion on medical aspects of the experiments and results. The manuscript was written under the lead of F.K., with contributions from M.B and D.P. All authors have given approval to the final version of the manuscript.

Funding:

This research received no funding.

Informed Consent Statement: Informed consent was obtained from all subjects involved in the study.”

Data Availability Statement: The experimental data used in this study are available from heIDATA, which is an institutional repository for the research data of Heidelberg University (DOI link will be provided at a later stage).

Acknowledgments: We kindly thank the Biogeochemistry group of the Institute of Earth Sciences at Heidelberg University for technical support of this study. Leonard Ernst for drawing Figure 6. We are grateful to John Hamilton for providing comments on an earlier version of the manuscript. This paper is dedicated to Prof. J.T.G Hamilton, who died in April 2022.

Conflicts of Interest: The authors declare no conflict of interest.

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