

1 **Climate and climate change in the Austrian–Swiss region of the**  
2 **European Alps during the twentieth century according to**  
3 **Feddema**

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11  
12 **Abstract**

13 Feddema's (Physical Geography 26: 442–466, 2005) generic climate classification method is  
14 applied to study climate and climate change in the Austrian–Swiss region of the European  
15 Alps during the course of the twentieth century. A fine-tuned version of it is also tested in  
16 addition to the original scheme. Monthly precipitation (P) and air temperature (T) data at a  
17 spatial resolution of 10'x10' are taken from the Climatic Research Unit (CRU) TS 1.2  
18 database to construct 30-year and 50-year period averages. It is shown that the alpine climate  
19 is sufficiently heterogeneous to make it unnecessary to perform fine-tuning of the original  
20 scheme for its characterization on the meso-β scale. It is also demonstrated that data  
21 organizational effects are much less intense than the effects caused by the fine-tuning. The  
22 area heterogeneity of climate and climate change types is the highest in the vicinity of lakes  
23 (Austria: Lake Constance; Switzerland: Lakes Geneva, Neuchâtel, Biel, Zurich and  
24 Constance) and along river valleys (Austria: the Danube, Drava, and Mur; Switzerland: the  
25 Aare and Ticino). The dominant climate change process was drying in Austria and warming  
26 in Switzerland. Large areas characterized by cold and saturated climate in the Central Eastern  
27 Alps did not experience climate change during the twentieth century.

## 28 **1 Introduction**

29 Generic climate classification methods serve the best characterization of climate in the given  
30 region. Among other climate classification methods, for instance, genetic methods or mostly  
31 mathematically established methods, these generic methods use as little data as possible and  
32 are the simplest, which enables their classroom applications while also being biophysically  
33 based. According to Essenwanger (2001) all these features are of crucial importance from the  
34 viewpoint of well-established climate classification.

35 Among generic methods the two most commonly used methods in the scientific  
36 community are the Köppen-Geiger (Geiger 1961) scheme and the Thornthwaite-based  
37 methods (Thornthwaite 1948; Feddema 2005). The main difference between them is in  
38 estimating the available heat. Köppen (1900, 1923, 1936) solved this problem as simply as  
39 possible using temperature data only. Thornthwaite (1948) recognized that potential  
40 evapotranspiration (PET) is a better indicator for this, so his method is PET-based not only in  
41 estimating thermal but moisture characteristics also. This PET-based method increases both  
42 the complexity and the applicability. Its applicability is increased since PET is a more suitable  
43 quantity for fine-tuning than temperature. Both methods are constructed for global scale  
44 analysis. There is an abundance of studies where the Köppen-Geiger (Geiger 1961) scheme  
45 was used for climate (Kottek et al. 2006; Peel et al. 2007) or climate change (Rubel and  
46 Kottek 2010) analyses. Thornthwaite's original method (Thornthwaite 1948) was never  
47 applied on the global scale, only its very simplified version constructed by Feddema (2005).  
48 There are also regional-scale analyses in which Köppen's (1936) scheme was also in the  
49 majority (e.g. Alvares et al. 2014; Engelbrecht and Engelbrecht 2016; Baker et al. 2010; Stern  
50 et al. 2000). Note that such analyses refer to countries or regions where the climate is fairly  
51 heterogeneous. Applications at even smaller scales are not common, but recently there have  
52 been such attempts (e.g. Ács et al. 2015; Breuer et al. 2016). These studies refer to a  
53 continental climate where the relief is quite uniform. To the best of our knowledge, there are  
54 no similar treatments referring to upland regions though the climate of upland regions is  
55 strongly heterogeneous even at local scales.

56 In this work, Feddema's (2005) method is used to characterize an upland climate,  
57 specifically, the climate of the European Alps, in its Austrian–Swiss region during the course  
58 of the twentieth century. We chose this region since it is located at about the same  
59 geographical latitude as Hungary. So, the climate differences between the lowland and the  
60 upland can be attributed mainly to differences in relief. The data needed are monthly  
61 precipitation (P) and temperature (T). They are taken from the Climatic Research Unit (CRU)  
62 TS 1.2 database (Mitchell et al. 2004). The analysis is performed using data of three 30-year  
63 periods (1901–1931; 1936–1965; 1971–2000) and two 50-year periods (1901–1950; 1951–  
64 2000).

65 We wish to answer the following questions: 1) Is it necessary to fine-tune Feddema's  
66 (2005) original scheme to be able to characterize the structure of climate on the mezo- $\beta$  scale  
67 (20–200 km)? 2) What are the main features of the alpine climate in the Austrian–Swiss  
68 region according to Feddema? 3) What are the main features of climate change in this region  
69 during the twentieth century? Finally 4) to what extent are the results sensitive to data  
70 organizational effects?

## 71 **2 Method and data**

72 **2.1 Feddema’s original scheme**

73 The original scheme is well presented in Feddema (2005). It is also discussed in Ács et al.  
 74 (2015), therefore only the most important features will be presented. Feddema’s (2005)  
 75 climate classification is based on calculation of PET and a moisture index  $I_m$ , which is defined  
 76 as follows:

$$77 \quad I_m = \begin{cases} 1 - \frac{PET}{P}, & \text{if } P > PET, \\ 0, & \text{if } P = PET, \\ \frac{P}{PET} - 1, & \text{if } P < PET. \end{cases}$$

78 PET is calculated after McKenney and Rosenberg (1993) since we did not find a special PET-  
 79 formula for alpine climate regions. The same parametrization works well (Ács et al. 2015) in  
 80 the Pannonian Plain, where the climate is continental. The method uses T and monthly mean  
 81 of potential sunshine duration as input.

83 **2.2 Feddema’s fine-tuned scheme**

84 The basic characteristics of Feddema’s fine-tuned scheme referring to the European Alps are  
 85 presented in Tables 1, 2 and 3.

86 **Table 1** Thermal types used in Feddema’s fine-tuned scheme version

<b>Thermal type</b>	<b>Annual PET (mm·year<sup>-1</sup>)</b>
cool	600–800
moderately cold	545–600
cold	300–545

87 **Table 2** Moisture types used in Feddema’s fine-tuned scheme version

<b>Moisture type</b>	<b>Moisture index (<math>I_m</math>)</b>
saturated	0.66–1.00
wet	0.47–0.66
moderately wet	0.33–0.47
moist	0.00–0.33
dry	-0.33–0.00

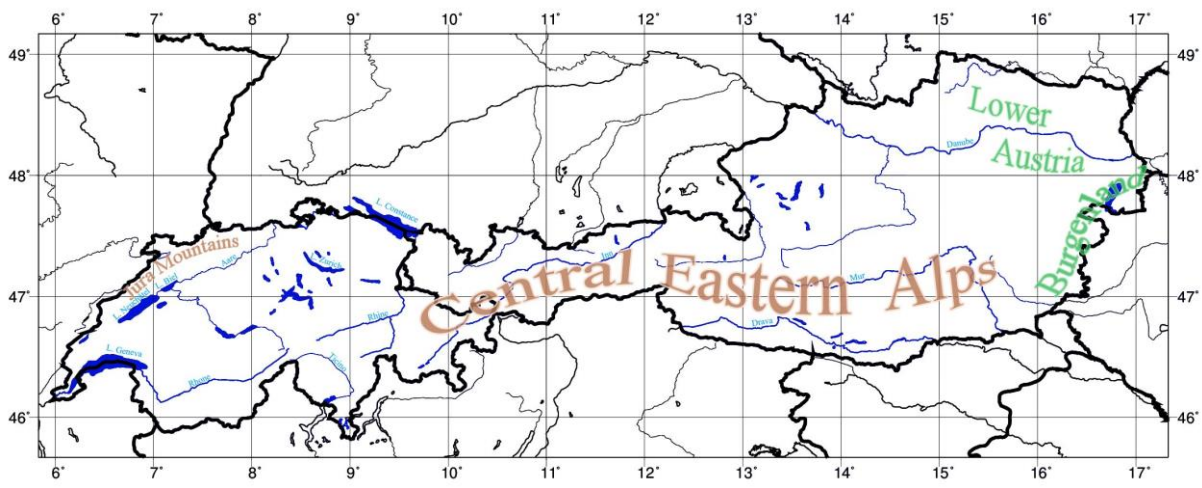
88 **Table 3** The magnitude of seasonal variability according to Feddema’s fine-tuned scheme  
 89 version

<b>Magnitude of seasonal variability</b>	<b>Annual range of <math>I_m</math></b>
low	0.00–0.50
medium	0.50–1.00
high	1.00–1.25
very high	1.25–1.50
extreme	1.50–2.00

90 Concerning thermal types, we introduced the “moderately cold” type in addition to the “cool”  
91 and “cold” categories. Concerning moisture types, the procedure was similar: the “moderately  
92 wet” category is used beside the categories “dry”, “moist”, “wet” and “saturated”. The criteria  
93 for deciding which climatic variable possesses seasonality remained unchanged. Nevertheless,  
94 the criteria for characterizing the magnitude of seasonal variability (Table 3) were slightly  
95 modified. We introduced the new category “very high” besides the categories “low”,  
96 “medium”, “high” and “extreme”.

## 97 2.3 Data

98 Monthly temperature and precipitation data are taken from the Climatic Research Unit (CRU)  
99 data center of the University of East Anglia. Data are part of the CRU TS 1.2 database  
100 (Mitchell et al. 2004), refer to the period 1901–2000 and possess a spatial resolution of  
101  $10' \times 10'$  (approximately  $18 \times 18 \text{ km}^2$ ). The studied region is presented in Fig. 1 together with  
102 the most important geographical designations used in the study. The area contains 1536 pixels  
103 located between the  $5.83^\circ$ – $17.33^\circ/45.67^\circ$ – $49.17^\circ$  longitude/latitude lines. In the analysis, we  
104 focused only on the Austrian–Swiss region using thirty-year (in total 71 fields) and fifty-year  
105 (51 fields in total) means of temperature and precipitation.



106

107 **Fig. 1** Austria and Switzerland in the studied region together with the major geographical  
108 designations used in the study

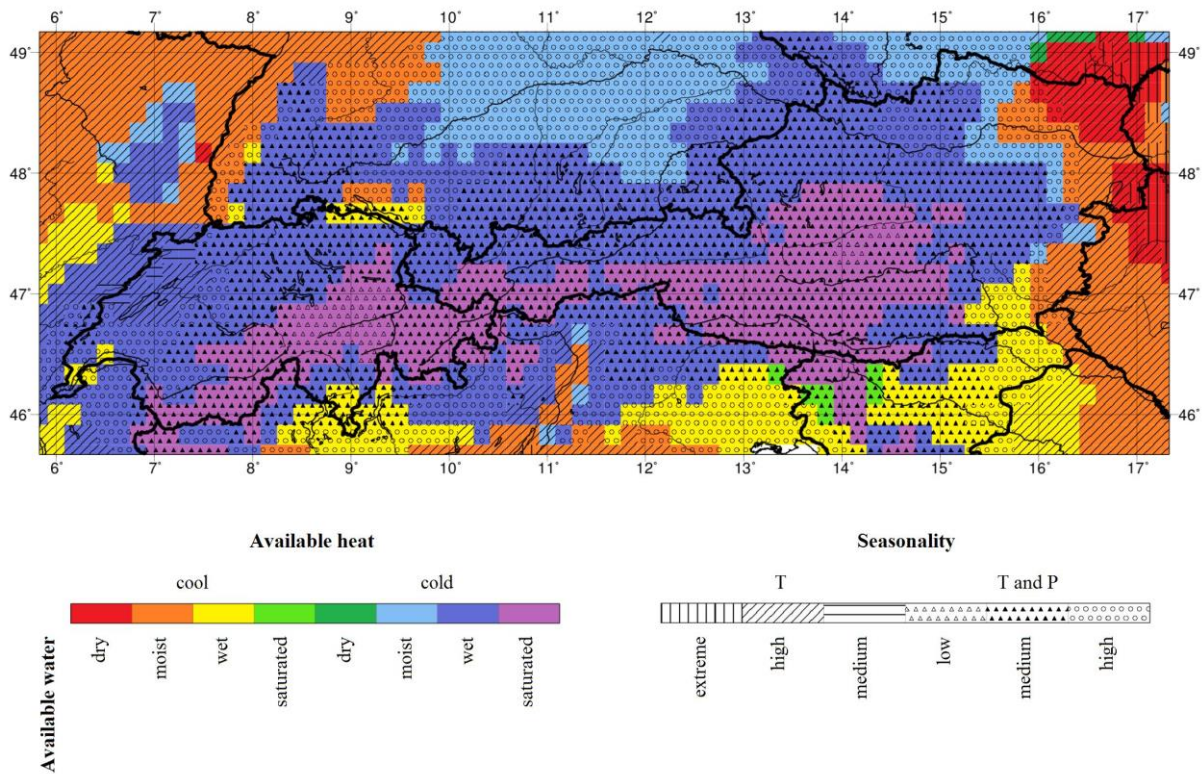
## 109 3 Results

110 The climate and the change of the climate in the Austrian–Swiss region of the European Alps  
111 during the course of the twentieth century will be discussed separately for 30-year and 50-  
112 year datasets. In the analysis, we mainly used Feddema’s (2005) original scheme. The fine-  
113 tuned model version applications based on 30-year datasets are used only for climate  
114 characterization purposes.

### 115 3.1 The climate of the Austrian–Swiss alpine region – original scheme

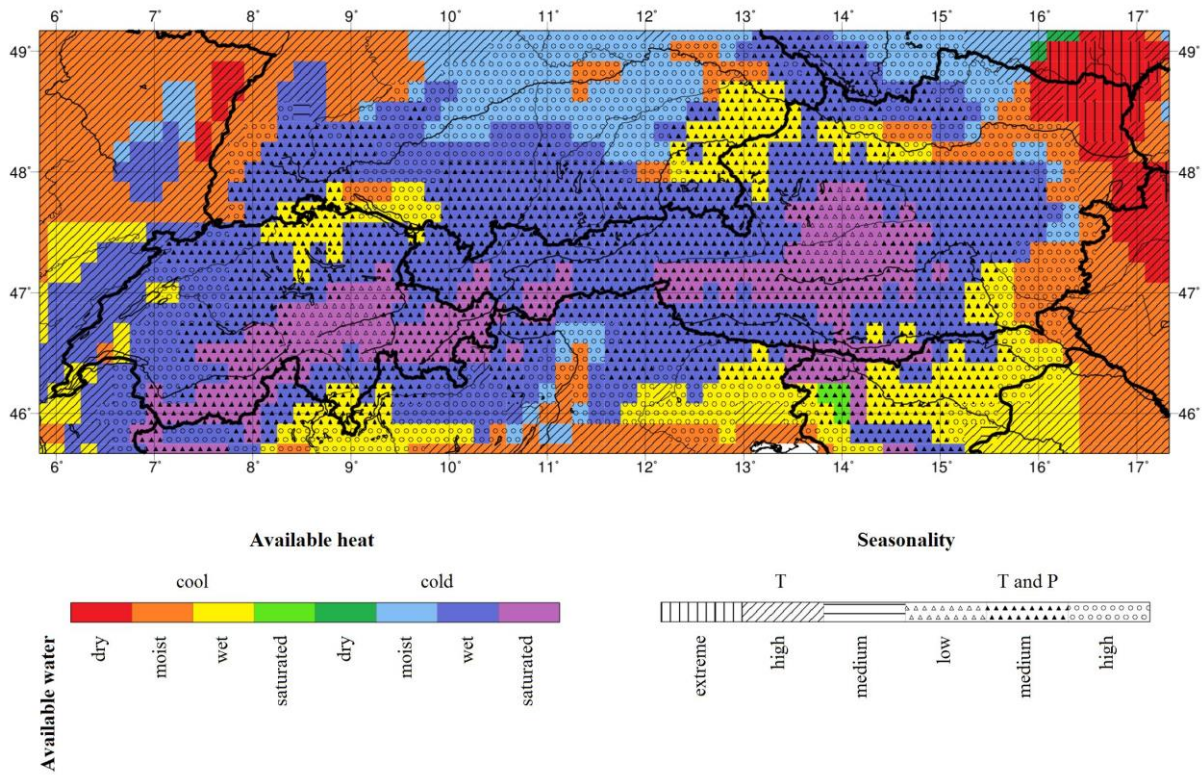
116 3.1.1 Thirty-year periods

117 Climate maps obtained from the original scheme at the beginning, in the middle and at the end  
118 of the twentieth century based on 30-year datasets are presented in Figs. 2, 3 and 4,  
119 respectively.



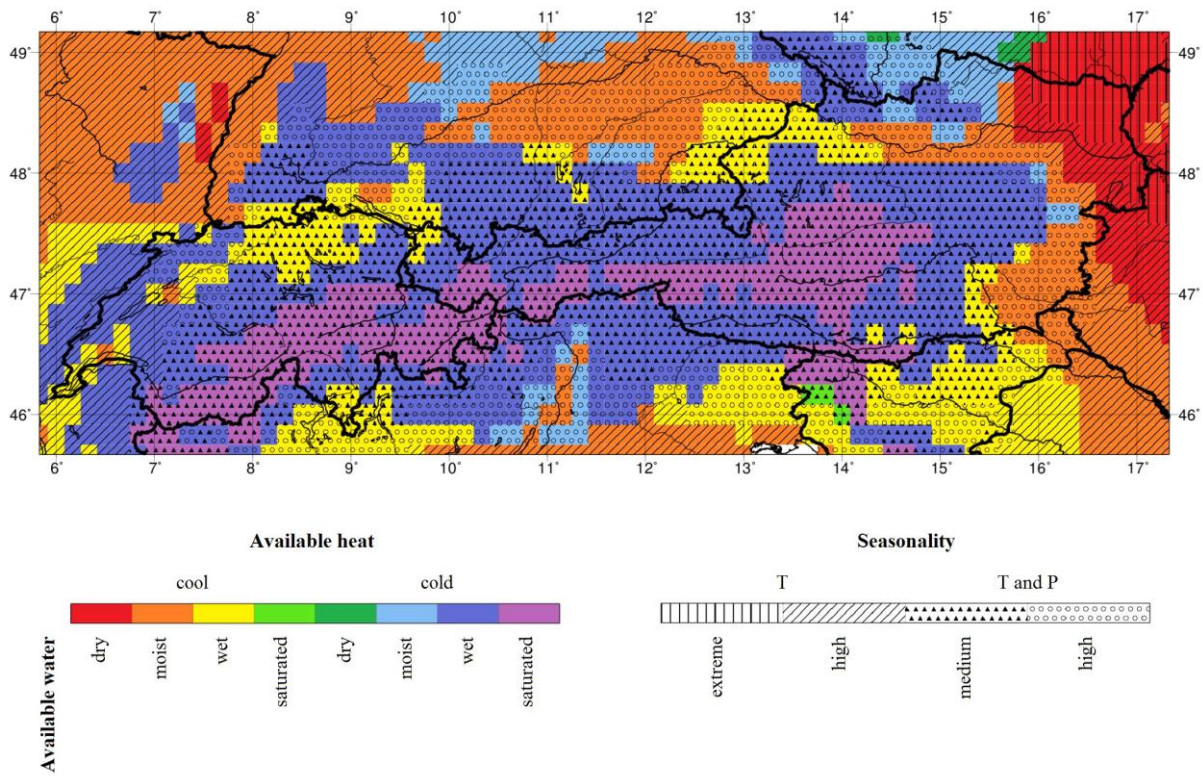
120

121 **Fig. 2** The climate of the Austrian–Swiss alpine region in the period 1901–1930 according to  
122 Feddema’s (2005) original scheme



123

124 **Fig. 3** The climate of the Austrian-Swiss alpine region in the period 1936-1965 according to  
 125 Feddema's original scheme



126

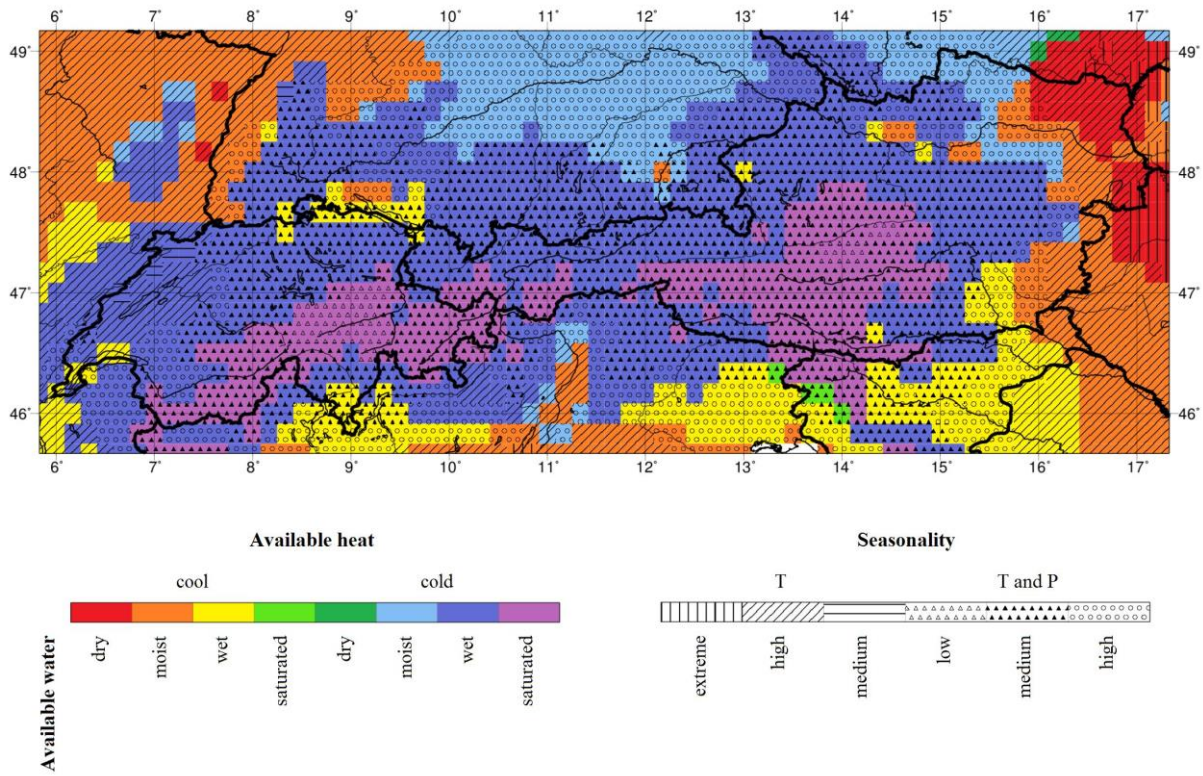
127 **Fig. 4** The climate of the Austrian–Swiss alpine region in the period 1971–2000 according to  
128 Feddema’s original scheme

129 The climate is the most heterogeneous in the north-eastern part of Austria. The greatest  
130 climate gradient is on the line that connects the eastern parts of Lower Austria and the Central  
131 Eastern Alps. The warmest and driest parts of the region are located in the Pannonian Plain  
132 (Lower Austria) and in the river and lake valleys. In these areas, 1) cool and dry, 2) cool and  
133 moist and 3) cool and wet climate types can be found with different seasonal characteristics.  
134 Extreme and high seasonality of T is a mainly Lower Austrian characteristic, but it can also be  
135 found on the shores of Lake Geneva, Lake Biel and along the Danube valley. Switzerland is  
136 cooler and wetter than Austria. In Switzerland, the climate type “cool and dry” does not exist.  
137 Areas with climate type “cool and moist” are rare, and can be found at Lake Geneva, Lake  
138 Biel and in the very small area of the Upper Rhine Plain. The climate type “cool and wet”  
139 covers larger areas than the climate type “cool and moist”. This can be found in the valleys of  
140 the Mur, Drava, Danube and Inn, on the shores of Lake Geneva, Lake Constance, in the  
141 region of the Swiss Plateau and in the Canton of Ticino, which is the most southern part of  
142 Switzerland. In these areas, seasonality is characterized by medium or high seasonality of  
143 both P and T. Note that this seasonal change characteristic is typical of alpine areas.

144 In the Central Eastern Alps, the “cold and wet” and “cold and saturated” climate types are  
145 prevailing with low or medium seasonality of both P and T. So, in total fifteen different  
146 climate types (different annual and seasonal characteristics) are registered by Feddema’s  
147 (2005) original scheme in the Austrian–Swiss region of the European Alps during the  
148 twentieth century.

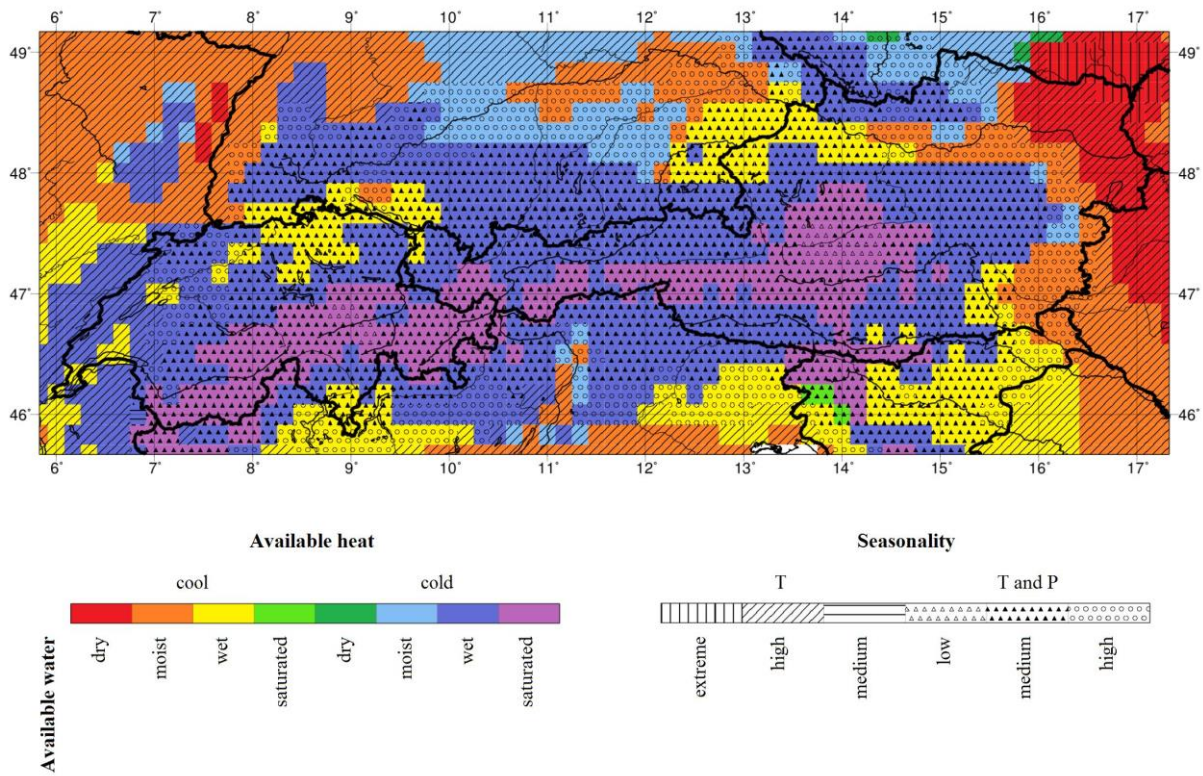
### 149 *3.1.2 Fifty-year periods*

150 The climate maps obtained using the original scheme referring to the 1<sup>st</sup> and 2<sup>nd</sup> part of the  
151 twentieth century based on 50-year datasets are presented in Figs. 5 and 6, respectively.



152

153 **Fig. 5** The climate of the Austrian–Swiss alpine region in the period 1901–1950 according to  
 154 Feddema’s original scheme



155

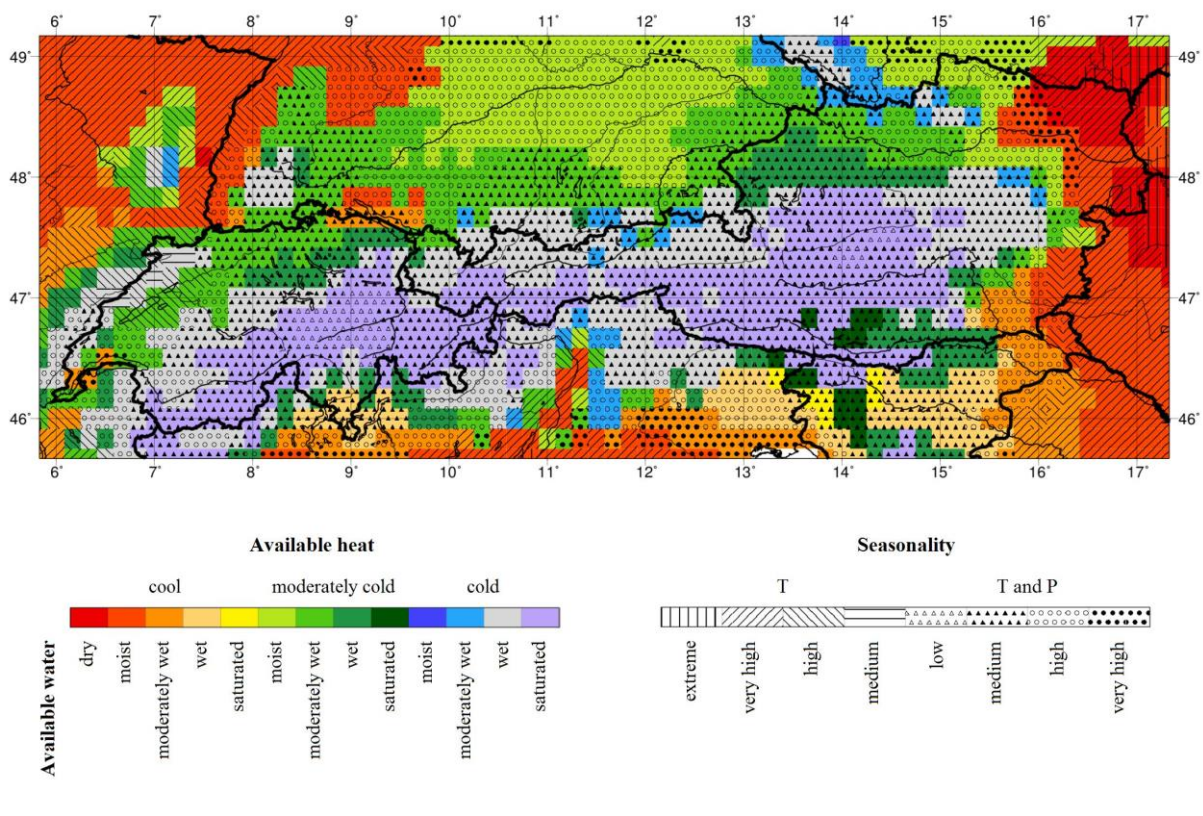
156 **Fig. 6** The climate of the Austrian–Swiss alpine region in the period 1951–2000 according to  
 157 Feddema’s original scheme

158 Inspecting Fig. 5, we see that it is very similar to Fig. 2. There are only minor differences  
 159 between them, that is, the spatial distribution of climate types is almost unaltered. Inspecting  
 160 Fig. 6, we easily detect that it is really very similar to Fig. 4. The climate types “cool and  
 161 moist” and “cool and wet” are also located in river valleys, on the shores of Lakes Geneva,  
 162 Neuchâtel, Biel and Zurich and in the Canton of Ticino. The climate types “cold and wet” and  
 163 “cold and saturated” are also typical only of alpine areas. As for 30-year periods, there are  
 164 fifteen climate types in total.

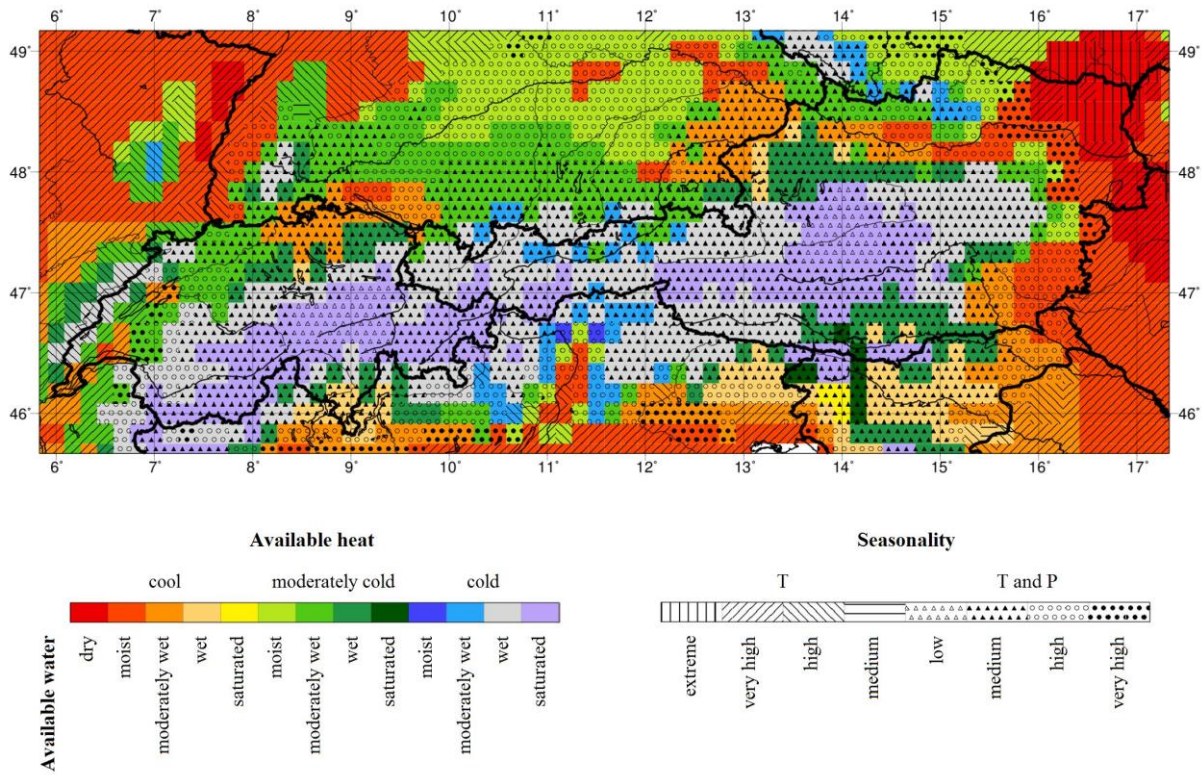
### 165 3.2 The climate of the Austrian–Swiss alpine region – fine tuned scheme

#### 166 3.2.1 Thirty-year periods

167 The climate maps obtained from the fine-tuned scheme at the beginning, in the middle and at  
 168 the end of the twentieth century based on 30-year datasets are presented in Figs. 7, 8 and 9,  
 169 respectively.

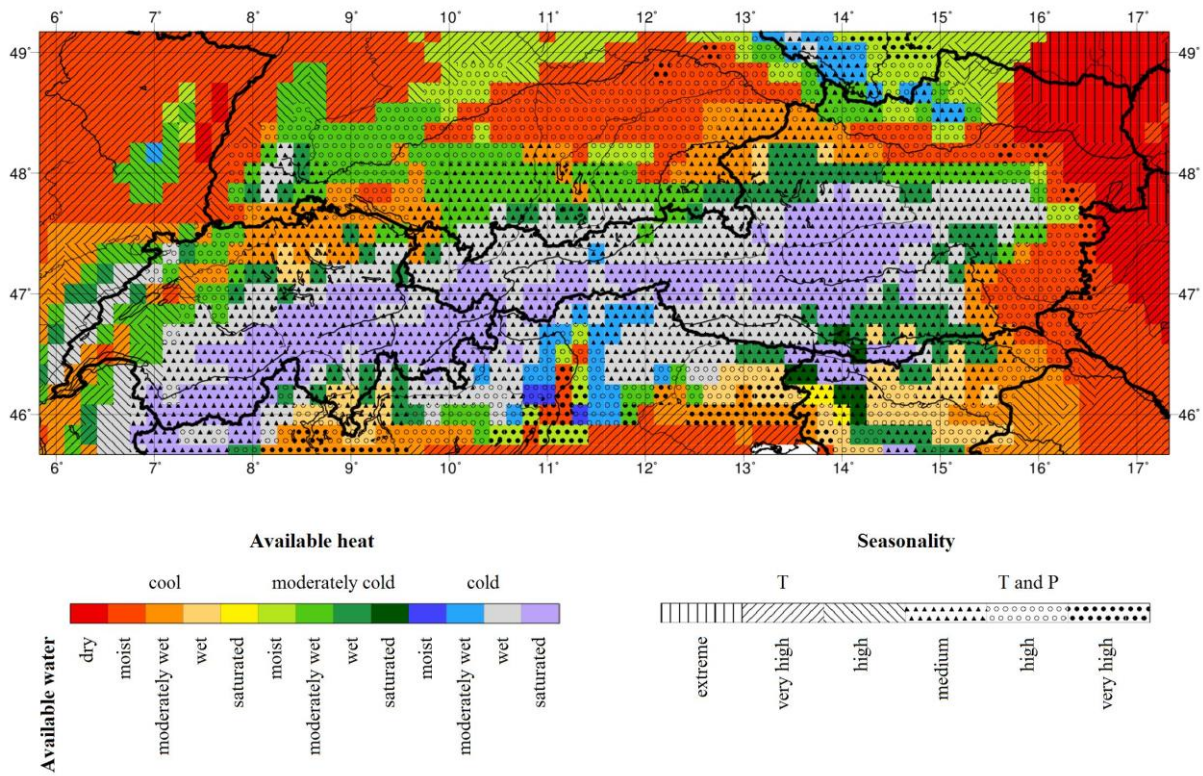


170  
 171 **Fig. 7** The climate of the Austrian–Swiss alpine region in the period 1901–1930 according to  
 172 Feddema’s fine-tuned scheme



173

174 **Fig. 8** The climate of the Austrian–Swiss alpine region in the period 1936–1965 according to  
 175 Feddema’s fine-tuned scheme



176

177 **Fig. 9** The climate of the Austrian–Swiss alpine region in the period 1971–2000 according to  
178 Feddema’s fine-tuned scheme

179 It is obvious that the number of climate types is much higher than in the cases (Figs. 2, 3 and  
180 4) when the original scheme was applied. Fine-tuning does not influence the area  
181 heterogeneity of climates in high mountainous regions, rather in river valleys and in the  
182 vicinity of large lakes. This is especially valid for the Mur and Danube in Austria, for Lake  
183 Geneva, the river Aare and the river Ticino in the most southern part of Switzerland. The  
184 change in climate types along the Mur valley is really pronounced (see Fig. 7). At the  
185 beginning of the twentieth century there were six climate types within a distance of about one  
186 hundred kilometers,:

- 187 - 1) cool and moderately wet with high variability of T and P,
- 188 - 2) moderately cold and moderately wet with high variability of T and P,
- 189 - 3) moderately cold and moderately wet with medium variability of T and P,
- 190 - 4) moderately cold and wet with medium variability of T and P,
- 191 - 5) cold and wet with medium variability of T and P and
- 192 - 6) cold and saturated with medium variability of T and P.

193

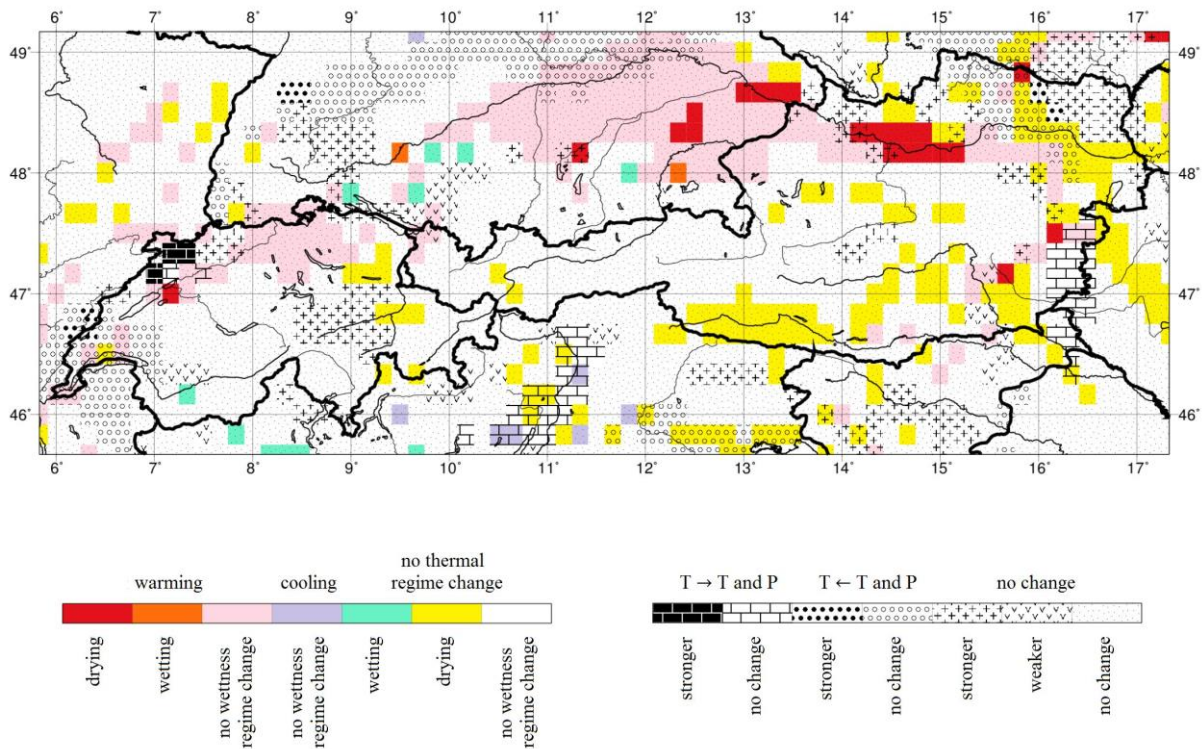
194 In the Danube valley the situation is similar. There were also six climate types within the  
195 distance of about two hundred kilometers at the beginning of the twentieth century (Fig. 7). In  
196 the Pannonian Plain region the climate is “cool and moist with very high variability of T”,  
197 while in the German–Austrian border region the climate type is “moderately cold, moderately  
198 wet with medium variability of T and P”. In the vicinity of Lake Geneva, the area  
199 heterogeneity of climate is also high. There are in total six climate types at the beginning of  
200 the twentieth century (Fig. 7). In the western parts of the lake the climate is “cool and  
201 moderately wet with high variability of T and P”, while in its eastern parts the climate is  
202 cooler and wetter. The greatest area heterogeneity of climates (high number of climate types  
203 in a small area) in Switzerland is in the Ticino valley. Four climate types can be found in this  
204 small region (Figs. 7, 8 and 9): “cool and wet”, “moderately cold and wet”, “cold and wet”  
205 and “cold and saturated”, while the magnitude of T and P variability is medium. Lastly: at the  
206 beginning of the twentieth century (Fig. 7) there were eighteen climate types in Austria and  
207 sixteen climate types in Switzerland. The total number of climate types in the Austrian–Swiss  
208 region is twenty-nine.

### 209 **3.3 Climate change in the region**

210 The climate change process is investigated only by using Feddema’s (2005) original scheme  
211 using both 30-year and 50-year datasets. In the first case, we compared climate maps referring  
212 to the beginning (Fig. 2) and to the end of the century (Fig. 4), while in the second case Figs.  
213 5 and 6 were compared. These comparisons were made as exactly as possible according to the  
214 procedure presented in the work of Breuer et al. (2016). The fine-tuned scheme is not applied  
215 for climate change analysis. It gave twenty-nine climate types, consequently there would also  
216 be an abundance of possible climate change types, which would enormously increase the  
217 complexity of analysis.

218 *3.3.1 Original scheme: thirty-year periods*

219 The spatial distribution of climate change types in the Austrian–Swiss alpine region during  
 220 the twentieth century obtained by Feddema’s (2005) original scheme and 30-year datasets is  
 221 presented in Fig. 10.



222  
 223 **Fig. 10** The climate change types in the Austrian–Swiss alpine region during the twentieth  
 224 century based on 30-year datasets according to Feddema’s original scheme

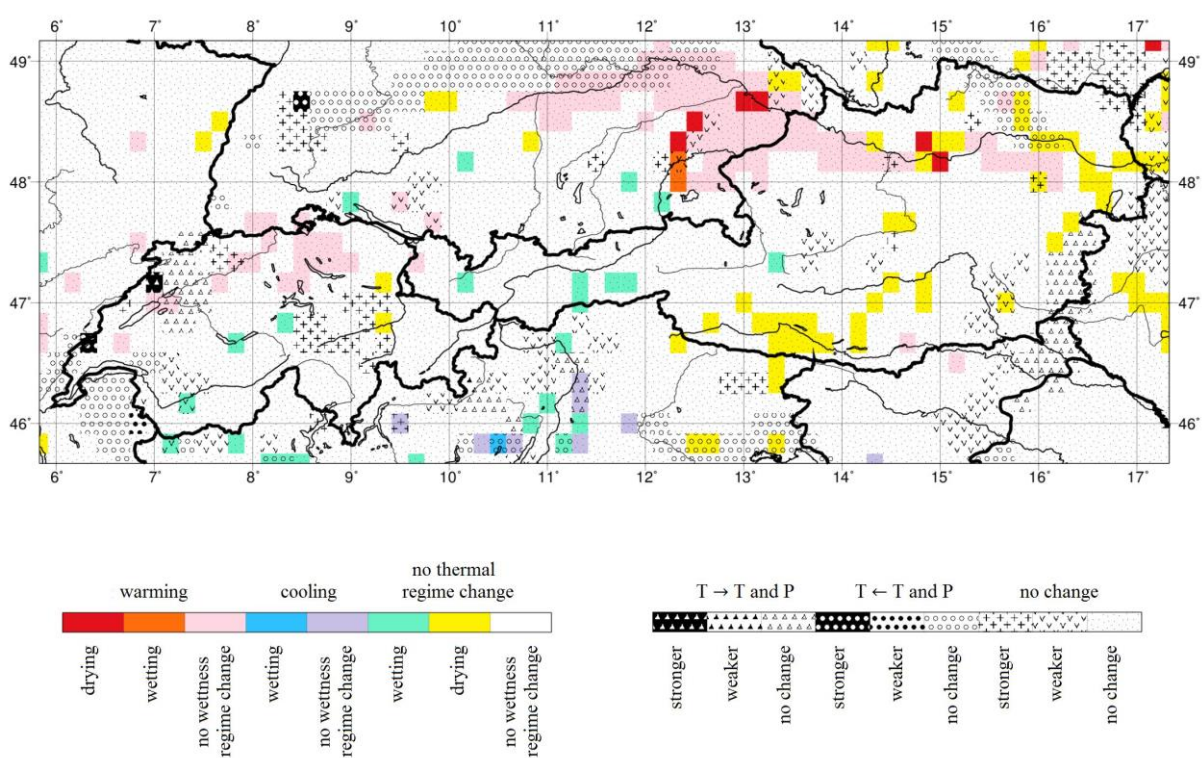
225 The climate change process was most pronounced along river valleys and around large lakes.  
 226 In Austria, the dominant process was drying. The most sensitive areas in Austria were the  
 227 valleys of the Danube, Drava and Mur and the shores of Lake Constance with the process of  
 228 drying even being observable in the Drava valley. The climate change process is more  
 229 complex along the Danube valley. Beside “single” changes of warming or drying, the  
 230 “double” change of warming and drying (in brief warming/drying) has also occurred. Only  
 231 “single” changes of warming or drying can be found in the Mur valley. Warming was  
 232 registered around Lake Constance. Changes in seasonal characteristics can be observed in  
 233 Burgenland and in northern parts of Lower Austria. In some areas, seasonal type changes  
 234 occurred: seasonality of T and P transformed into seasonality of T (T and P → T) and, vice  
 235 versa, seasonality of T transformed into seasonality of T and P (T → T and P). There are also  
 236 areas (for instance, in the vicinity of Vienna) where the magnitude of variability increased.

237 The dominant climate change process in Switzerland is warming. The “single” change  
 238 process of warming can be observed on the shores of Lake Geneva, in the river Aare valley  
 239 and in the region between the river Aare and Lake Zurich. The “double” change process of  
 240 warming and drying is not typical for Switzerland, only one such pixel can be found around

241 Lake Biel. The “single” change process of drying can also be observed, for instance, at Lake  
 242 Geneva, or in the Rhine valley. Interestingly, there is also one pixel where the process of  
 243 wetting occurred. This can be found in the Rhône valley. Changes in seasonality can also be  
 244 observed. Such changes are typical for Lake Geneva and the Jura Mountains, but some such  
 245 areas are also located in the Central Eastern Alps. Note that there is no climate change in the  
 246 area of the Canton of Ticino.

247 *3.3.2 Original scheme: fifty-year periods*

248 The spatial distribution of climate change types in the Austrian–Swiss alpine region during  
 249 the twentieth century obtained by Feddema’s (2005) original scheme and 50-year datasets is  
 250 presented in Fig. 11.



251  
 252 **Fig. 11** The climate change types in the Austrian–Swiss alpine region during the twentieth  
 253 century based on 50-year datasets according to Feddema’s original scheme

254 This climate change map is similar to the map presented in Fig. 10 but there are also  
 255 significant differences. Beside the “single” climate change processes of warming or drying,  
 256 the process of wetting is also represented not only in Switzerland but also in Austria. In  
 257 Austria, the processes of warming and drying are approximately equally represented and the  
 258 areas where the “double” climate change process of warming and drying is represented are  
 259 considerably reduced. There are only two such pixels in the Danube valley. Otherwise, the  
 260 Danube valley is the area where climate change is the most pronounced. In the Drava valley,  
 261 the process of drying is typical as in the former case. There is also a drying process in the Mur  
 262 valley but not so pronounced. There, changes in seasonality could also be observed.

263 In Switzerland, warming is the dominant “single” climate change process. Areas  
264 characterized by warming can be found around Lakes Zurich, Neuchâtel and Biel and in the  
265 river Aare valley. Changes in seasonality can be observed approximately equally around Lake  
266 Geneva, in the Jura Mountains and in the Central Eastern Alps. These changes are various,  
267 not only the type but also the magnitude of variability changed. Concerning magnitude of  
268 variability, both the increase (around the Rhine) and decrease (around the Rhône) can be  
269 observed. As in the former case, no climate change was registered in the area of Canton of  
270 Ticino.

## 271 **4 Conclusion**

272 Concluding remarks will be organized according to the questions given in the last paragraph  
273 of section 1. 1) The fine-tuning of Feddema’s (2005) scheme has larger effect on the  
274 appearance of climate maps than the organization of data. The fine-tuning increased the  
275 number of climate types from fifteen to twenty-nine, but with this the crowding of the map  
276 increased and transparency decreased. Based on this, we do not suggest that a fine-tuned  
277 version of Feddema’s (2005) scheme be used in upland regions, like the European Alps, for  
278 climate classification purposes on the mezo- $\beta$  scale. 2) The area heterogeneity of climate and  
279 climate change types is higher around lakes, along river valleys and lower in upland regions.  
280 These high heterogeneity regions are located in the valleys of the Danube, Mur and Drava in  
281 Austria, and in the valleys of the Aare and Ticino as well as around Lakes Geneva, Neuchâtel,  
282 Biel, Zurich and Constance in Switzerland. In Austria, the dominant climate change process  
283 was drying, while in Switzerland it was warming. The highest climate type heterogeneity area  
284 can be found in the Canton of Ticino, at the same time there was no climate change registered  
285 during the course of the twentieth century. It is to be noted that there are extended areas of  
286 cold and saturated climate in the Central Eastern Alps. In these areas no climate change  
287 occurred during the twentieth century.

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